



Microcosm experiments to control anaerobic redox conditions when studying the fate of organic micropollutants in aquifer material

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ABSTRACT

The natural processes occurring in subsurface environments have proven to effectively remove a number of organic pollutants from water. The predominant redox conditions revealed to be one of the controlling factors. However, in the case of organic micropollutants the knowledge on this potential redox-dependent behavior is still limited. Motivated by managed aquifer recharge practices microcosm experiments involving aquifer material, settings potentially feasible in field applications, and organic micropollutants at environmental concentrations were carried out. Different anaerobic redox conditions were promoted and sustained in each set of microcosms by adding adequate quantities of electron donors and acceptors. Whereas denitrification and sulfate-reducing conditions are easily achieved and maintained, Fe- and Mn-reduction are strongly constrained by the slower dissolution of the solid phases commonly present in aquifers. The thorough description and numerical modeling of the evolution of the experiments, including major and trace solutes and dissolution/precipitation of solid phases, have been proven necessary to the understanding of the processes and closing the mass balance. As an example of micropollutant results, the ubiquitous beta-blocker atenolol is completely removed in the experiments, the removal occurring faster under more advanced redox conditions. This suggests that aquifers constitute a potentially efficient alternative water treatment for atenolol, especially if adequate redox conditions are promoted during recharge and long enough residence times are ensured.

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1. Introduction

The ultimate motivation of this work is artificial recharge of aquifers. Artificial recharge is beneficial both in quantitative (augmentation of groundwater resources, long term underground storage, etc.) and qualitative terms (overall improvement of water quality during aquifer passage: decreasing of

suspended solids, pathogens, nitrogen, phosphates, metals and dissolved organic carbon). The interest in this technique is also related to the capability of subsoil processes to partially or totally remove organic contaminants from water (Aronson et al., 1999 and references therein; Christensen et al., 2001 and references therein; Neuhauser et al., 2009). Nowadays a great scientific effort is dedicated to assess whether organic micropollutants could also be effectively removed (Barber et al., 2009; Díaz-Cruz and Barceló, 2008 and references therein; Heberer, 2007 and references therein; Hoppe-Jones et al., 2010). A number of such compounds are not eliminated by conventional water treatments (Gros et al., 2010; Onesios

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et al., 2009 and references therein; Petrovic et al., 2009 and references therein; Stackelberg et al., 2007). The passage of water through the soil–aquifer system during artificial recharge may represent an alternative or complementary treatment for their removal.

As a general rule, the fate of organic pollutants within the aquifer depends on lithology, hydraulic and textural properties of the soil, temperature, physico-chemical properties of the specific compound, and microbial environment. Among all factors, the predominant redox state of the aquifer revealed to play a significant role (Aronson et al., 1999 and references therein; Bosma et al., 1996; Broholm and Arvin, 2000 and references therein; Christensen et al., 2001; Kao et al., 2003 and references therein). Since certain pollutants could be preferably removed under some particular redox conditions, such conditions could eventually be promoted in artificial recharge practices. Even more important, if different compounds are degraded under different redox environments, a water mass undergoing a sequence of redox states should have most of its initial contaminants eliminated.

Yet, in the case of organic micropollutants, the knowledge on a potential redox-dependent behavior is still limited. Beside field evidences (Drewes et al., 2003; Heberer et al., 2008; Montgomery-Brown et al., 2003; Pavelic et al., 2005; Tubau et al., 2010; and references therein), laboratory tests under specific and controlled simplified conditions have been carried out. However, many of the experiments reported in literature adopted settings that are not representative or directly applicable to aquifer systems.

Specifically, aerobic/anoxic biodegradability and transformation mechanisms of organic micropollutants have been investigated in model systems for wastewater treatment. That is, laboratory experiments have been typically performed in wastewater matrices, and by using sludges from sewage treatment plants as adapted inocula (Clara et al., 2004; Quintana et al., 2005; Stasinakis et al., 2009; Zwiener et al., 2002). Their fate under aerobic/anaerobic or specific redox conditions have also been largely studied in aquatic environments. In these cases river-bed sediments rich in organic materials have been incubated with river water (Davis et al., 2006; Löffler et al., 2005; Radke et al., 2009) or with solutions/culture media containing specific electron acceptors (Bradley et al., 2001; Crawford et al., 1998; Lu et al., 2009; Somsamak et al., 2001). Some tests involved bacterial isolates, and have been carried out using standard silica sand or sintered materials as solid matrix for the colonization of the microorganisms (Crawford et al., 2000; Katz et al., 2001; Stucki et al., 1995). Finally, not only in the aforementioned studies but also when soil and aquifer material were included (Krueger et al., 1998; Schulz et al., 2008; Ying et al., 2008), the experiments have been often performed with concentrations of the target compounds from hundreds of $\mu\text{g L}^{-1}$ to tens of mg L^{-1} .

The above works are indeed useful to demonstrate the susceptibility of specific micropollutants to microbial or abiotic transformation, to understand degradation pathways, and to identify intermediate or stable metabolites. However, the organic content of aquifer materials, which may influence sorption and partitioning behavior of organic micropollutants, could be lower, the potential development of a sequence of redox states and the removal of micropollutants depends on the local native microorganisms, and target

pollutants are found at concentrations some order of magnitudes lower.

Finally, quite limited laboratory experiments, a number of them related to managed aquifer recharge practices, resemble real subsurface environments (Baumgarten et al., 2011; Hua et al., 2003; Mansell and Drewes, 2004; Massmann et al., 2008; Rauch-Williams et al., 2010; Scheytt et al., 2004). In such experiments, the fate of organic micropollutants has been usually assessed within the range of redox conditions developing naturally in the system and being representative of those actually occurring at field site, namely aerobic, anoxic (prevailing denitrifying) and seldom unspecified anaerobic conditions. The identification of potential abiotic processes by performing analogous sterile experiments was not always included in such studies. Therefore, the potential effect of various redox states (especially the most reducing ones) on the fate of a wide range of organic micropollutants in subsurface environments still remains to be investigated.

In this context, the aim of our work was to create and sustain diverse anaerobic redox conditions in systems involving natural aquifer material and settings potentially feasible in artificial recharge sites, and to study in such environments the behavior of organic micropollutants at realistic concentrations.

The present paper describes thoroughly the experimental methodology and settings adopted. Details on the selection of the type/quantities of electron donors and acceptors used to stimulate the desired redox conditions have been integrated. Limited information on the design criteria is usually provided in studies on the fate of organic pollutants adopting this approach (Bosma et al., 1996; Bradley et al., 2001; van der Zaan et al., 2009; Weiner et al., 1998; Ying et al., 2008). This hinders running analogous studies with some different setting (substrates, electron acceptors, durations, etc.).

The description of the microcosms' hydrochemical evolution is also presented. Often this is not/poorly monitored or only incipiently reported in laboratory studies on the fate of organic contaminants, especially when focused on their transformation pathways, nor the actual occurrence of the expected/stimulated redox condition is verified (Baumgarten et al., 2011; Bosma et al., 1996; Bradley et al., 2001; Gröning et al., 2007; Krueger et al., 1998; Rauch-Williams et al., 2010; Schulz et al., 2008; van der Zaan et al., 2009; Weiner et al., 1998; Ying et al., 2008). We conjecture that the assessment of the geochemical state, which could be quite complex as in natural subsurface environments, and its quantitative numerical modeling has to be included in this type of studies for a more complete interpretation of the experimental results, and for the potential subsequent design of real field applications.

Finally, as example of application of the study to the fate of organic micropollutants in different redox environments, the results for the ubiquitous but still barely investigated β -blocker atenolol are presented in the paper.

2. Materials and methods

2.1. Sediments, water and micropollutants

The experimental set up included various sets of microcosms, each microcosm consisting of natural sediments and synthetic water spiked with a mixture of organic micropollutants.

Sediments were obtained from a test site for artificial recharge of groundwater located in Sant Vicenç dels Horts (Barcelona, Spain). The aquifer consists of quaternary alluvial sediments, mainly gravel and sand with a small fraction of lutites. Sediment samples for the experiments were collected prior to the starting up of the facilities, from a pit excavated in the bottom of the infiltration pond, namely from an oxic unsaturated horizon at about 1 m depth. They were sieved through a 1 mm grid to remove the coarse fraction, which was expected to be less active for surface and microbially mediated reactions. The sieved sediments were immediately used for assembling the microcosms or stored for a maximum of 2 days at 25 °C inside aluminum paper.

Chemical and mineralogical characteristics of the sediments used in the experiments are summarized in Table S1 of the Supporting Information. X-ray diffraction (XRD) of powdered samples was used in an attempt to identify the minerals present in the sediment. Analysis was performed with a Bruker D-5005 diffractometer. Results were obtained using Cu radiation, with secondary Graphite monochromator. The analytical conditions were: $\theta/2\theta$ geometry, collecting data in the range (2θ) between 4° and 60° with a step scan of 0.05°, 3 s per step measuring time. The evaluation of the spectra was made by using the DiffraSuite™ software and identification of chemical compounds by means of the PDF database Release 2001, Data Sets 1–51 plus 70–89.

Total nitrogen, total carbon and organic carbon content were analyzed using an organic elemental analyzer with on line combustion–reduction–gas chromatography (TCD detector) model EA series 1108 (Thermo Fisher Scientific), set at conditions within the provider-recommended range. Data acquisition and calculations were done with the Eager 200 software (Thermo Fisher Scientific).

The grain size distribution was measured with the Laser diffraction particle-size analyser Coulter LS230 (Beckman Coulter Inc., Fullerton, CA, USA) with a Detection Limit of 0.04 μm . The content in Mn and Fe(III) associated to oxide-hydroxides and oxides was obtained by sequential extraction (steps 1 to 4 from Dold, 2003).

Water used for the preparation of all the experiments, called “common water” in the following, was artificially prepared to mimic the recharge water at the test site (from the Llobregat river water) except for the organic carbon content, which at this stage was set equal to zero. Its theoretical composition is shown in Table S2 of the Supporting Information.

The mixture of organic micropollutants used in all experiments included drugs (atenolol, carbamazepine, diclofenac, gemfibrozil, ibuprofen, sulfamethoxazole), pesticides (atrazine, simazine, terbutylazine, prometryn, diuron, chlorphenvinfos, chlorpyrifos, diazinon), estrogens (estrone, β -estradiol), PAHs (naphthalene, acenaphthene, fluorene, anthracene, Phenanthrene, benz[a]anthracene, chrysene, pyrene, fluoranthene, benzo[b]fluoranthene, benzo[k]fluoranthene, benzo[a]pyrene, dibenzo[a]anthracene, indeno[1,2,3-cd]pyrene and benzo[ghi]perylene), surfactant degradation products (4-tert-octylphenol, 4-nonylphenol), a phthalate (bis-diethylhexyl phthalate) and a biocide (triclosan). The selection of the compounds was based on the micropollutants occurrence in the Llobregat river (Céspedes et al., 2005; Muñoz et al., 2009; Rodríguez-Mozaz et al., 2004).

High purity (>96%) analytical standards of atenolol, carbamazepine, diclofenac, gemfibrozil, ibuprofen, simazine, diuron and estrone, and of the isotopic analog atenolol d7 used as surrogate standard for quantification of atenolol were supplied by Sigma-Aldrich. The standard containing the 16 PAHs at a concentration of 2000 mg L^{-1} in dichloromethane:benzene (1:1) as well as high purity (>96%) analytical standards of all the remaining compounds were purchased from AccuStandar. Individual stock solutions were prepared in methanol or in an appropriate solvent according to their properties. Working standard mixtures were then prepared at different concentrations by dilution of the individual stock solutions in methanol, and were used to prepare the spiking solution for the experiments (resulting concentration in the “initial water” described in Section 2.4 was 10 $\mu\text{g L}^{-1}$ for 4-octylphenol and 4-nonylphenol, and 1 $\mu\text{g L}^{-1}$ for the rest of compounds) and to prepare the aqueous calibration standards (concentration range 1–1500 ng L^{-1} , surrogate standard 200 ng L^{-1}). Stock and working standard solutions were stored at –20 °C in the dark.

2.2. Biotic experiments – creating sustainable redox conditions

A different anaerobic redox state was promoted in four different sets of batches by stimulating one specific step of the natural redox sequence for organic matter degradation (Table 1). To this end, easily degradable organic compounds were provided as electron donors and, depending on the target redox condition, NO_3^- , Mn(III/IV), Fe(III) or SO_4^{2-} was added as specific electron acceptor.

Nitrate and sulfate were incorporated to the “common water” by dissolving magnesium nitrate hexahydrate and sodium sulfate, respectively. Mn(III/IV) and Fe(III) oxide-hydroxides were incorporated to the sediments as finely ground natural psilomelane and mixed ferrihydrite/goethite (1:10 in weight), respectively.

Sodium acetate and the methanol used as solvent in the micropollutants spiking solution were adopted as easily degradable substrates. They were incorporated to the “initial water” (Section 2.4) by dissolving anhydrous sodium acetate and when spiking the micropollutants mixture, respectively. In fact, the organic micropollutants introduced represented potential electron donors too, but their concentrations were so low that their effect on redox condition build-up was expected to be minimal.

The selection of the type of substrate was based on a revision of the existing literature and on preliminary scoping experiments (results not shown) regarding the degradation feasibility and rate for different organic compounds. Ideally, the selected substrate should promote the build-up of the desired redox conditions after a short lag-phase and form a small number of intermediate compounds (possibly being not fermentable) to facilitate the assessment of the chemical evolution of the system.

The total amounts of organic substrate and controlling electron acceptor were selected so as to reach the desired redox state and to sustain it during a significant lapse of time. This implies on one hand that the total amount of organic substrate had to be large enough to consume electron acceptors with reactions energetically more favorable than the target one.

Table 1

Sequence of the overall redox reactions for the microorganisms mediated degradation of organic matter (i.e. methanol and acetate ions in the present experiments). Biomass growth is ignored in the stoichiometries.

1a) $\text{CH}_3\text{OH} + 1.5 \text{O}_2 \rightarrow \text{HCO}_3^- + \text{H}^+ + \text{H}_2\text{O}$ 1b) $\text{CH}_3\text{COO}^- + 2\text{O}_2 \rightarrow 2\text{HCO}_3^- + \text{H}^+$	Aerobic respiration	Redox potential	Energy yield
2a) $\text{CH}_3\text{OH} + 1.2 \text{NO}_3^- + 0.2 \text{H}^+ \rightarrow \text{HCO}_3^- + 0.6 \text{N}_2 + 1.6 \text{H}_2\text{O}$ 2b) $\text{CH}_3\text{COO}^- + 1.6 \text{NO}_3^- + 0.6 \text{H}^+ \rightarrow 2\text{HCO}_3^- + 0.8 \text{N}_2 + 0.8 \text{H}_2\text{O}$	Nitrate reduction	+	↑
3.1a) $\text{CH}_3\text{OH} + 3 \text{MnO}_2(\text{s}) + 5 \text{H}^+ \rightarrow \text{HCO}_3^- + 3 \text{Mn}^{2+} + 4 \text{H}_2\text{O}$ 3.1b) $\text{CH}_3\text{COO}^- + 4 \text{MnO}_2(\text{s}) + 7 \text{H}^+ \rightarrow 2\text{HCO}_3^- + 4 \text{Mn}^{2+} + 4 \text{H}_2\text{O}$ 3.2a) $\text{CH}_3\text{OH} + 6 \text{MnOOH}(\text{s}) + 11 \text{H}^+ \rightarrow \text{HCO}_3^- + 6 \text{Mn}^{2+} + 10 \text{H}_2\text{O}$ 3.2b) $\text{CH}_3\text{COO}^- + 8 \text{MnOOH}(\text{s}) + 15 \text{H}^+ \rightarrow 2\text{HCO}_3^- + 8 \text{Mn}^{2+} + 12 \text{H}_2\text{O}$	Mn oxide reduction		↑ increasing
4a) $\text{CH}_3\text{OH} + 6 \text{Fe}(\text{OH})_3(\text{s}) + 11 \text{H}^+ \rightarrow \text{HCO}_3^- + 6 \text{Fe}^{2+} + 16 \text{H}_2\text{O}$ 4b) $\text{CH}_3\text{COO}^- + 8 \text{Fe}(\text{OH})_3(\text{s}) + 15 \text{H}^+ \rightarrow 2\text{HCO}_3^- + 8 \text{Fe}^{2+} + 20 \text{H}_2\text{O}$	Fe ox/hydroxide reduction		
5a) $\text{CH}_3\text{OH} + 0.75 \text{SO}_4^{2-} \rightarrow \text{HCO}_3^- + 0.75 \text{HS}^- + 0.25 \text{H}^+ + \text{H}_2\text{O}$ 5b) $\text{CH}_3\text{COO}^- + \text{SO}_4^{2-} \rightarrow 2\text{HCO}_3^- + \text{HS}^-$	Sulphate reduction	-	↓

On the other hand, for each potential selection of organic substrates the total amount of controlling electron acceptor initially available had to (slightly) exceed the stoichiometric quantity necessary for their complete mineralization. We used the stoichiometries in which biomass formation is not considered (see Table 1), since after the original sources of organic carbon have been depleted, dead cells could be recycled and degraded coupling with the reduction of some electron acceptor.

Further details on the selection of the initial amounts of electron donors and acceptors can be found in the Supporting Information (Text S1).

The definitive concentrations of the organic substrates (and the corresponding amount of target electron acceptor) to be initially available in each experiment could be established according to the degradation rates observed in the preliminary rough tests performed when selecting the type of substrates, where their potential toxicity toward microorganisms could be excluded too, and according to the desired duration of the experiments.

Namely, in the present tests the design initial concentration of methanol was fixed by the quantity of spiking solution of micropollutants added to the “initial water”, i.e. 2.7 mmol L^{-1} . Regarding acetate, according to the design constraints the selected initial concentrations exceeded natural levels in aquifer system. Still, they were inside the range of concentrations already used in injection experiments and bioremediation scenario in subsurface environments (Baker et al., 1999; Kerkhof et al., 2011), i.e. within a range applicable to a potential stimulation of some specific redox condition in artificial aquifer recharge field sites.

In a complementary set of batches called “natural conditions”, with oxygen initially present in the system, neither additional electron acceptors were added to the “common water” or to the sediments, nor was sodium acetate added as electron donor to the “initial water”. That is, no specific redox state was deliberately stimulated and the organic matter degradation reactions were expected to develop sequentially (Table 1, set “a” of reactions), until complete depletion of either the electron donors or acceptors available in the system. Also in this experiment the design initial concentration of methanol was 2.7 mmol L^{-1} , fixed by the quantity of spiking solutions of micropollutants added to the “initial water” during the assembling procedure.

The initial concentrations of electron donors and electron acceptors present in the five biotic experiments are shown in Table 2. Notice that in the NO_3^- -reducing experiment, due to some unidentified problem during the assembling procedure, some additional 2.9 mmol L^{-1} of methanol turned out to be present apart from the 2.7 mmol L^{-1} proceeding from the spiking solution, resulting in an actual initial methanol concentration higher than expected.

2.3. Abiotic experiment

A sterile experiment was also conducted as common control reference for the biotic tests to identify potential abiotic processes affecting the micropollutants. The absence of biodegradation processes precludes the evolution of the redox sequence from evolving. Therefore, no information would be gained by repeating the abiotic experiment for each specific electron acceptor. They would never be used.

Table 2
Initial analytical concentration of electron donors and acceptors in the five sets of microcosms.

Type of experiment	Electron donors		Electron Acceptors			
	Initial analytical DOC (mM)	Contribution of the easily degradable organic substrates to the Initial analytical DOC [%]	Initial measured $O_2(ac)$ [mM]	Initial analytical NO_3^- [mM]	Initial analytical SO_4^{2-} [mM]	Initial Mn (III/IV) Initial Fe (III)
NO_3^- -reducing	9.7	CH_3COO^- 42.79 CH_3OH 57.21	0.0	6.7	2.0	The amount originally present in the sediments
Mn (III/IV)-reducing	6.7	CH_3COO^- 60.50 CH_3OH 39.50	0.0	0.1	2.3	0.4 g of natural psilomelane added to the original sediment
Fe (III)-reducing	7.8	CH_3COO^- 68.17 CH_3OH 31.83	0.0	0.1	2.3	The amount originally present in the sediments
SO_4^{2-} -reducing	10.2	CH_3COO^- 75.38 CH_3OH 24.62	0.0	0.1	5.3	The amount originally present in the sediments
Natural conditions	2.5	CH_3COO^- 0.00 CH_3OH 100.00	0.2	0.1	2.3	The amount originally present in the sediments

The amount originally present in the sediments
0.95 g of natural mixed ferrihydrite/goethite (1:10 in weight) added to the original sediment

The synthetic water used to prepare this experiment (“abiotic water”) consisted in the previously described “common water” plus the same additional amounts of magnesium nitrate hexahydrate and sodium sulfate used respectively in the NO_3^- -reducing and SO_4^{2-} -reducing experiments. Further on, during the assembling procedure (Section 2.4), the same amount of acetate used in the SO_4^{2-} -reducing experiment was added to the sterile “initial water”. The initial concentration of methanol was 2.7 mmol L^{-1} , fixed by the quantity of spiking solution of micropollutants added. Thus, in the end, the abiotic experiment was characterized by the maximum NO_3^- , SO_4^{2-} , Na, Mg and DOC concentration existing among the 5 biotic experiments.

Details on the sterilization procedure are given in Section 2.4.1.

2.4. Experimental procedure

2.4.1. Assembling the microcosms

The collected sediments were air dried at laboratory temperature (25°C), homogenized in steel containers, and distributed in fractions of 120 g (air-dry weight) into 0.3 L glass bottles. The previously defined amounts of Mn(III/IV) or Fe(III) oxide-hydroxides powder were mixed with the sediments of each bottle for the Mn(III/IV)- and Fe(III)-reducing experiments, respectively. The 0.3 L glass bottles were then placed inside a glove box under Ar atmosphere (maximum 0.1% of O_2).

The “common water” was prepared in a glass amber bottle. The previously defined amounts of NO_3^- or SO_4^{2-} were added in the case of the NO_3^- -reducing and SO_4^{2-} -reducing experiments, respectively. The water was bubbled with Ar (purity $\geq 99.999\%$) during about 1 hour to remove all oxygen from water and bottle headspace. Afterwards, the bottle was closed with a screw-cap plus a PTFE protection seal and placed into the glove box under Ar atmosphere, where the remaining part of the assembling procedure was performed.

The water was finalized by adding the predefined amount of sodium acetate and the spiking solution of micropollutants. After sampling the resulting “initial water” for chemical analyses, 0.24 L of it were added to each one of the 0.3 L glass bottles already containing the sediments. The assembling procedure was concluded by closing the bottles with screw-caps plus a PTFE protection seal, and gently shaking. A remaining headspace of about 15 mL was left in each bottle.

The bottles were removed from the glove box and enveloped with aluminum foil to prevent photodegradation. Then, they were incubated under controlled temperature ($25 \pm 2^\circ\text{C}$) and gently shaken few times during their lifetime (once every 2 days during the first week; once a week during the rest of the first month; then, once every 30 to 45 days) as well as the day before being sacrificed.

The “natural conditions” experiment was conducted without Ar bubbling or assembling within a glove box. Instead, dissolved oxygen was allowed in the water and oxygen gas was initially present in the headspace of the bottles.

In the case of the abiotic experiment, prior to the beginning of the assembling procedure, the sediments and the “abiotic water” were sterilized three times (once a day in three consecutive days) with autoclave at $T = 121^\circ\text{C}$ and $P = P_{atm} + 1 \text{ atm}$ for 20 min; moreover, the glove box was sterilized with UV light before entering the material. As an additional precaution,

0.22 mmol L⁻¹ of mercury chloride were added (as microbial poison) to the “initial water”.

2.4.2. Disassembling the microcosms

Duplicate bottles were sacrificed at each sampling time according to predefined sampling schedules (Table S3 of the Supporting Information). These had been defined according to the expected degradation rates of the organic substrates and micropollutants reported in the literature and those estimated in the preliminary scoping experiments (Section 2.2). Some sampling times were set equal for the different experiments, to facilitate comparisons.

One at a time, the two bottles were opened under Ar atmosphere, chemical parameters were measured, and aqueous samples for general chemistry and micropollutants analysis were collected and stored according with each laboratory recommendations.

The sterility of the abiotic experiment was verified six times along its duration. An aliquot of water from devoted microcosms was spread on tryptic soy agar (TSA) plates and incubated in duplicate at 25 °C under aerobic conditions (for 1 week) and anaerobic conditions (for 2 weeks). None of the plates demonstrated microorganism growth.

2.5. Monitoring and analysis

Samples collected for analyzing Cl⁻, NO₃⁻, NO₂⁻, SO₄²⁻, PO₄³⁻, F⁻, NH₄⁺, DOC and COD (Chemical Oxygen Demand) were filtered through 0.45 µm PALL Acrodisc® Sterile Syringe Filters with Supor® membrane and frozen. Anions were analyzed by ion chromatography using an ICS-1000 instrument. The analytical error was estimated to be 14% for PO₄³⁻ and 13% for the remaining anions. NH₄⁺ concentration was analyzed with a selective electrode Orion 9512. DOC was analyzed by 680 °C combustion catalytic oxidation/NDIR method using a TOC-V CSH instrument. The estimated analytical error was 20%. COD was analyzed by colorimetry with the spectrophotometer Spectroquant Nova 60.

Samples for the analysis of Fe and Mn, Ca, Mg, Na, K and minor elements were also filtered at 0.45 µm, acidified and stored at 4 °C. They were later analyzed by inductively coupled plasma atomic emission spectrometry (ICP-AES) using a Thermo Jarrel-Ash Iris Advantage HS instrument. Detection limits were 100 µg L⁻¹ for K and Na, and 50 µg L⁻¹ for the rest. The analytical error was estimated below 3%. In the ICP-AES analyses, calibration with three laboratory sets of standards was performed every 10 samples, and regression coefficients of the calibration curves exceeded 0.999.

pH and temperature (Thermo Scientific 9157BN Triode pH electrode, refillable), electrical conductivity (Hanna Instruments, 76302W conductivity probe) and dissolved oxygen (Hanna Instruments, HI 76407/4 DO probe) were measured during the assembling/disassembling procedure with specific electrodes. Alkalinity was measured with a drop test kit Taylor K-1726, with a precision of 0.5 mmol L⁻¹.

Samples for analysis of atenolol were filtered at 0.45 µm using WATERS Syringe filter with PTFE membrane. Then, they were kept frozen until analysis, which was performed by using on-line solid phase extraction–liquid chromatography–tandem mass spectrometry. Briefly, water samples (10 mL) spiked with the isotopically labeled compound at a

concentration of 200 ng L⁻¹, were extracted with the aid of an automated on-line SPE sample processor Prospekt-2 from Spark Holland (Emmen, The Netherlands) connected in series with the LC–MS/MS instrument. Sample preconcentration was performed by passing 5 mL of the sample through a previously conditioned (1 mL MeOH plus 1 mL HPLC water) Oasis HLB Prospekt™ cartridge (10 × 1 mm) from Waters (Mildford, MA, USA). After sample loading, the cartridge was washed with 1 mL of a 5% methanol water solution and further eluted with the chromatographic mobile phase. Chromatographic separation was performed with a Binary HPLC pump Model 1525 from Waters using a Purospher STAR RP-18e column (125 × 2 mm, 5 µm particle diameter, from Merck, Darmstadt, Germany) and gradient elution with methanol and water as mobile phase. MS/MS detection was performed in the selected reaction monitoring (SRM) mode acquiring 2 SRM transitions per compound and 1 SRM transition per surrogate using a TQD triple–quadrupole mass spectrometer from Waters equipped with an electrospray interface. Quantitation was performed by the internal standard method using the corresponding deuterated compound as surrogate standard. Due to defective functioning (inaccurate sample volume acquisition) of the SPE processor, the first 3 results of the Mn(IV)-, Fe(III)- and SO₄-reducing experiments and the first 2 results of the “natural conditions” experiment could only be considered as semiquantitative.

2.6. Modeling

The hydrogeochemical evolution of the experiments was simulated by using CHEPROO (Bea et al., 2009), a Fortran 90 module using object-oriented concepts that simulates complex hydrobiogeochemical processes. The thermodynamic database used was that of EQ3NR code (Wolery, 1992).

The precipitation of calcite, magnesite, dolomite, siderite and amorphous iron sulfide was assumed to be controlled by kinetics. A simplified formulation was used to describe their reaction rate r [mol m⁻³ s⁻¹]:

$$r = k\sigma(1 - \Omega) \quad (1)$$

where k is a rate constant [mol m⁻² s⁻¹], σ is the reactive surface of the mineral [m² m⁻³], and Ω is the saturation ratio [–].

The microbially mediated redox reactions for the organic substrates degradation (only the easily degradable substrates were considered, i.e. acetate and methanol) were described by kinetic rate laws based on Monod expressions:

$$r_i = k_i S \cdot \frac{TEA}{K_{i,TEA} + TEA} \cdot \frac{K_{inhib,1}}{K_{inhib,1} + I} (\Omega_i - 1)^{\eta_i} \quad (2)$$

where r_i is the rate of consumption of the substrate [mol L⁻¹ s⁻¹], k_i is the first order rate coefficient [s⁻¹], S is the substrate concentration [mol L⁻¹], TEA is the concentration of the particular Terminal Electron Acceptor [mol L⁻¹], $K_{i,TEA}$ is the Monod half saturation constant with respect to TEA [mol L⁻¹], I is the concentration of an inhibiting substance (e.g. a competing TEA) [mol L⁻¹] and $K_{inhib,1}$ is the inhibition constant [mol L⁻¹]. The last term of expression (2), also called far-from-equilibrium term and where Ω_i is the saturation ratio (ratio between ion activity product and equilibrium constant

of the redox reaction “i”) and η_i is an experimental parameter, describes the thermodynamic constraint for the redox reaction “i”. In our case, Ω_i is $\ll 1$ because of the high concentrations of organic substrates and the precipitation as mineral of large amounts of reaction products. Thus, the effect of such term on the absolute value of r_i is insignificant. Multiple Monod and inhibition terms could be included in Eq. (2) if deemed necessary.

3. Results and discussion

3.1. General water chemistry

According to the reactions of Table 1, a general trend of decrease in DOC and increase in alkalinity was expected in the biotic experiments. A simultaneous decrease in the concentration of the target dissolved electron acceptors was expected for the NO_3^- and SO_4^{2-} -reducing experiments, whereas an increase in the concentration of Fe(II) and Mn(II), products of the reduction of the target solid electron acceptors, was expected in the Mn- and Fe-reducing experiments. Details on the geochemical evolution (experimental datasets and simulations) of the biotic experiments are given below. The results from duplicate batches showed a satisfactory reproducibility at all sampling times. Actually, when plotting the measurements plus the error bars from each batch, there was always some overlap. Thus, the following graphics report the average of results and manual measurements from the duplicate bottles.

It is worth mentioning that in the following, for the sake of simplicity and according to the near neutral pH characterizing the experiments, alkalinity and Dissolved Inorganic Carbon (DIC) have been considered practically equal to the HCO_3^- concentration.

Regarding the abiotic experiment, the hydrochemistry remained practically constant for the whole time as expected (results not shown).

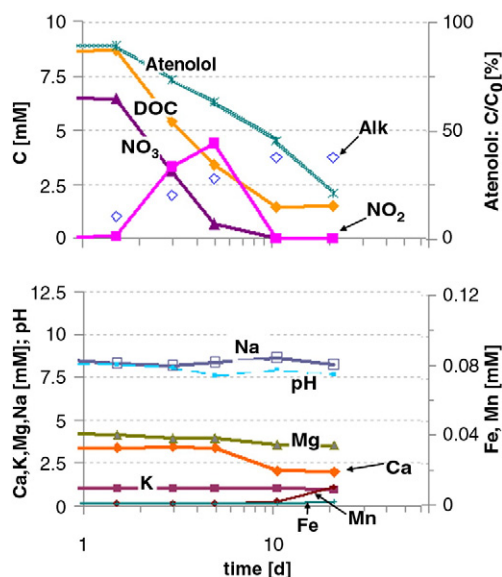


Fig. 1. Chemical evolution with time in the NO_3^- -reducing experiment.

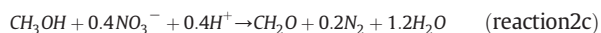
3.1.1. NO_3^- -reducing experiment

Results from the NO_3^- -reducing experiment are shown in Fig. 1. During the first 10 days, DOC decreased from 9.7 mmol L^{-1} to 1.5 mmol L^{-1} . Afterwards it remained practically constant. At day 10, the 6.7 mmol L^{-1} of nitrate initially present in the water have disappeared. Nitrite concentration began to increase after only some 12 h, reaching a maximum at day 5 and becoming completely depleted by day 10. Very low concentrations of dissolved manganese and iron were detected after day 10, presumably from the dissolution and reduction of small quantities of the Mn and Fe oxides naturally present in the sediment. Sulfate remained constant during the whole experiment.

These observations suggest that nitrate reducing conditions were established within a short period (~ 0.5 days) of microbial adaptation and dominated the system during the first 10 days. The increase of nitrite, followed by its depletion, reflected the actual pathway for nitrate reduction, with nitrite being an intermediate product between nitrate and nitrogen. After day 10, a different more reducing condition was established.

The experiment was planned to guarantee complete depletion of organic carbon with excess of nitrate, allowing nitrate-reducing conditions to dominate during longer time. However, the actual initial nitrate and DOC concentrations (6.7 and 9.7 mmol L^{-1} , respectively) turned out to be different from their designed amounts (7.4 and 6.9 mmol L^{-1} , respectively) due to some unidentified problems during the assembling of the experiment. Consequently, nitrate and nitrite were completely depleted, while some organic carbon was still present after day 10.

The decrease of DOC (8.3 mmol L^{-1}) up to day 10 exceeded its expected stoichiometric removal (6.5 mmol L^{-1}) calculated by assuming that the only process that can change nitrate concentration was reduction coupled with organic matter oxidation, and according to reactions 2a and 2b of Table 1 in which biomass formation is not taken into account. This suggests that some organic carbon was used into microorganisms' growth. An estimation of such investment could be made by introducing an additional reaction. In fact, since carbon in bulk biomass has a redox state of 0, the formation of biomass requires a partial oxidation in the case of methanol (redox state of carbon = -2). Coupling it with the reduction of nitrate, the following stoichiometry could be written:



where CH_2O has been used as simplified formula for biomass.

By using reactions 2a and 2b of Table 1, and reaction 2c, a conversion of 2.2 mmol L^{-1} of organic carbon into biomass could be finally calculated, i.e. about 27% of the total organic carbon consumption.

Alkalinity increased with time, but its final value at day 10 (3.75 mmol L^{-1}) was smaller than the expected one (7 mmol L^{-1}) calculated by taking into account the amount of organic substrate mineralized under the previous hypothesis. Part of this gap could be explained by the net reduction of Ca and Mg concentrations (0.9 mmol L^{-1} and 1.1 mmol L^{-1} , respectively), which suggests that precipitation of CaCO_3 , MgCO_3 or mixed carbonates was limiting the actual increase

in bicarbonate concentration. The Saturation Index (S.I.) of these minerals during the experiment supports this hypothesis. It ranged between 0.24 and 0.71 for calcite, and from 0.06 to 0.45 for magnesite. Throughout the paper, S.I. values were calculated using the PHREEQC code with WATEQ thermodynamic database (Parkhurst and Appelo, 1999). After the conclusion of the experiment, inspection of the sediment samples by SEM-EDS showed indeed small crystals of calcite and Mg–Ca carbonates on the surface of the sediment grains (Fig. 2A).

We next consider the equilibrium of the aqueous carbonate species with the gas in the headspace of the bottles. By day 10, about 0.3 mmol L^{-1} of inorganic carbon (representing the 7% of the total inorganic carbon inventory) have been transferred to the gas phase as $\text{CO}_2(\text{g})$. Finally, taking into account the precision of alkalinity measurements ($\pm 0.5 \text{ mmol L}^{-1}$), the overall inorganic carbon mass balance could be closed with an error of about 15%.

The simulations carried out with CHEPROO (Fig. 3) support the feasibility of the previous hypotheses. The most important parameters used are reported in Table S4 of the Supporting Information.

3.1.2. Mn(III/IV)-reducing experiment

Results from the manganese reducing experiment are shown in Fig. 4. DOC decreased with time from 6.7 to 1 mmol L^{-1} , starting after day 7 and reaching a significant removal rate after day

14. The small initial NO_3 (0.1 mmol L^{-1}) had already disappeared by day 7 (not shown), having only oxidized a small amount of DOC (a maximum of 0.2 mmol L^{-1}). Consistently with DOC decrease, dissolved Mn increased from day 7 reaching a concentration of 0.1 mmol L^{-1} at day 25, which is then maintained for the rest of the experiment. No Fe was detected and SO_4 concentration remained almost constant during the whole experiment.

Alkalinity increased slightly until day 25. Thereafter, up to day 42 it dropped down to a value that remained steady throughout the rest of the experiment. The net reduction of Ca and Mg concentrations (1.3 mmol L^{-1} and 1 mmol L^{-1} , respectively) and a lower than expected increase of Mn concentration suggested precipitation of Mg–Ca carbonates and MnCO_3 , limiting alkalinity and dissolved Mn. The computed S.I. with respect to calcite and rhodochrosite during the experiment (between -0.11 and 0.97 for calcite, and between 1.15 and 1.73 for rhodochrosite) supports this hypothesis. SEM-EDS examination of the sediments from the disassembled batches showed the presence of small crystals of calcite, Mn-bearing carbonates, Mg–Ca carbonates, and rhodochrosite (Fig. 2B).

Regarding the low Mn concentration detected, aside from the fact that a fraction of the original DOC was invested in biomass growth implying a smaller total Mn^{2+} production than that corresponding to the complete mineralization of all substrates, an additional explanation could be some Mn^{2+} adsorption on the

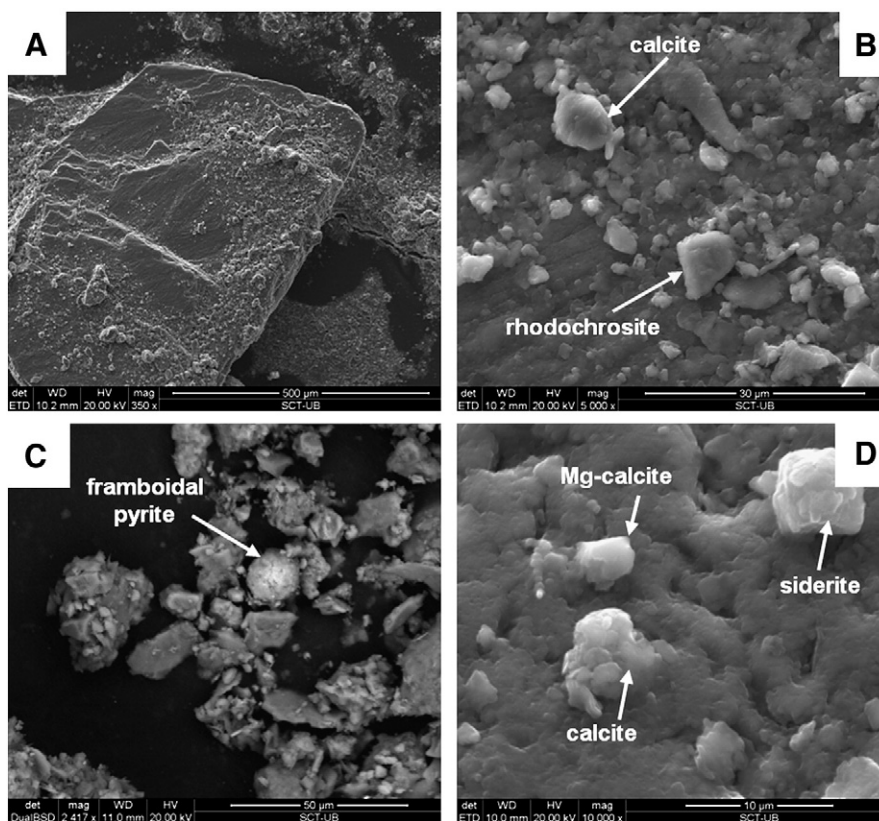


Fig. 2. SEM images of sediment samples from the disassembled batches. A) Neo-formed carbonate grains on the cleavage surface of a feldspar crystal present in the sediment in the NO_3 -reducing experiment. B) Precipitates of calcite and rhodochrosite in the Mn-reducing experiment. C) Precipitates of Ca-, Mg-Ca- and Fe-carbonates in the Fe-reducing experiment. D) Framboidal pyrite was occasionally observed in the Fe-reducing experiment, likely originated by the turnover of FeS previously precipitated.

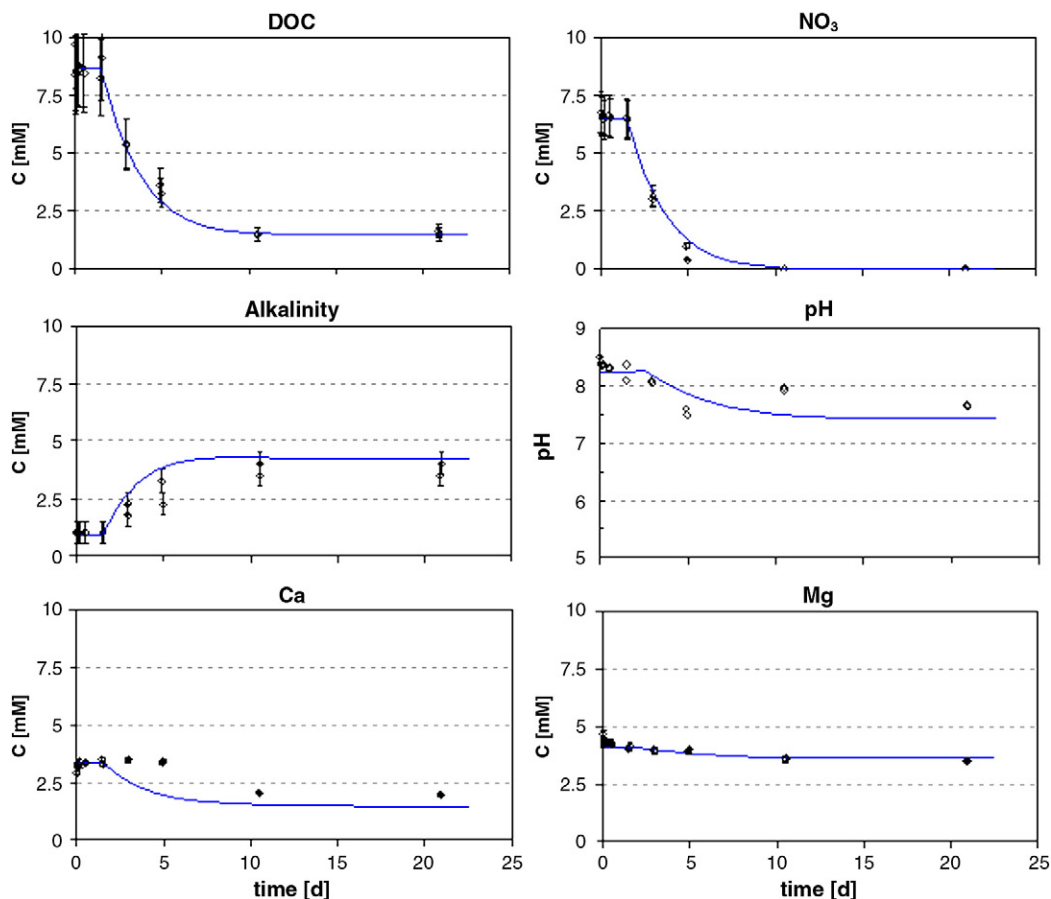


Fig. 3. Chemical evolution with time in the NO_3 -reducing experiment: simulations (solid lines) versus experimental data (dots).

clay surfaces (von Gunten and Zobrist, 1993; Wang and Van Cappellen, 1996).

In summary, DOC and the redox sensitive species suggest that Mn-reducing conditions were established after ~1 week of microbial adaptation and were then maintained during the test. The slow dissolution of the natural source of Mn(III/IV) used in

the experiment was likely representing a rate limiting factor for the Mn reduction. Since the exact identity of the Mn oxide-hydroxides added could not be confirmed, an unequivocal mass balance for C and Mn (according to reactions 3.1a to 3.2b of Table 1) could not be carried out and quantitative modeling was not performed.

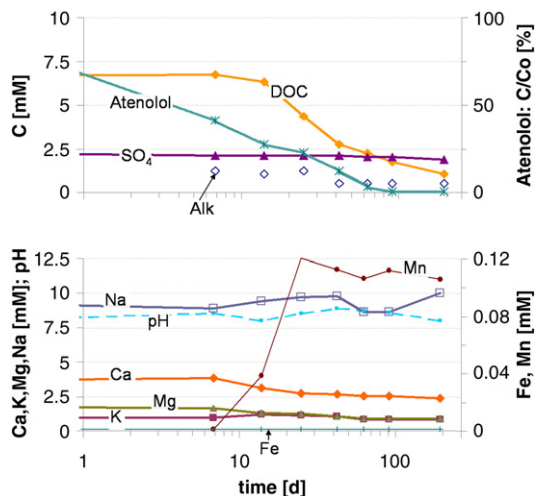


Fig. 4. Chemical evolution with time in the Mn(III/IV)-reducing experiment.

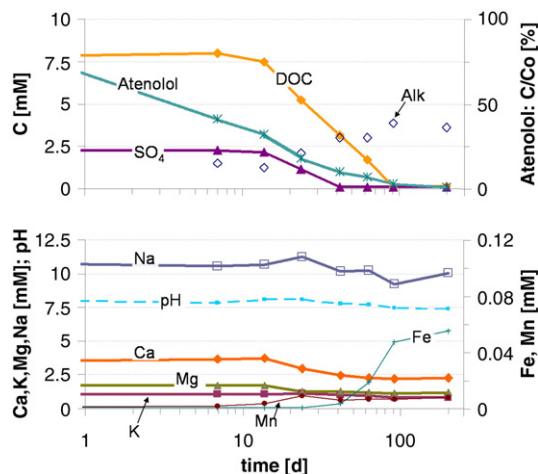


Fig. 5. Chemical evolution with time in the Fe(III)-reducing experiment.

3.1.3. Fe(III)-reducing experiment

Results from the iron reducing experiment are shown in Fig. 5. A small decrease of DOC could be observed starting from day 7. By then, the small initial NO_3^- (0.1 mmol L^{-1}) had already disappeared (not shown), having oxidized a small amount of DOC (a maximum of 0.2 mmol L^{-1}). Significant changes in water chemistry could be observed after day 14, but Fe was not detected until day 42, when the 2.3 mmol L^{-1} of initial SO_4 had been completely depleted. After day 42 the dissolved Fe (Fe(II) at the pH range of this experiment) increased slightly to about 0.06 mmol L^{-1} . DOC decreased with time from 7.8 to 3.2 mmol L^{-1} at day 42 and to complete depletion after day 91. Alkalinity increased from 1.2 to 3.0 mmol L^{-1} at day 42 and to 3.6 mmol L^{-1} thereafter. Ca and Mg concentrations decreased along the experiment: 1 and 0.6 mmol L^{-1} up to day 42, and 0.2 and 0 mmol L^{-1} after day 42, respectively. A very low concentration of dissolved Mn, never exceeding 0.01 mmol L^{-1} , was detected starting day 14. Probably it was produced by the reductive dissolution of some Mn mineral, naturally present in the sediments, causing a negligible effect on DOC concentration.

The above observations suggest that, after approximately 1 week of microbial adaptation, both SO_4 and Fe(III) were reduced until day 42 causing FeS precipitation, whose low solubility prevented build up of dissolved Fe. This was

confirmed by the dark color of the sediments at disassembling. Occasionally some framboidal pyrite has been observed by SEM-EDS on the surface of sediment grains (Fig. 2C), likely generated by the aging of precipitated FeS. Since dissolved Fe did not increase until SO_4 was exhausted, the rate of Fe(III)-reduction needs to be slower than SO_4 -reduction. While this would contradict the sequence of Table 1, it was not entirely surprising since the Fe(III) source was a natural solid phase. Slow dissolution of this source may be the rate limiting mechanism for Fe reduction. This means that HS^- could be in part accumulated in solution. Concomitant Fe(III)- and SO_4 -reduction and iron sulfide precipitation has already been observed in field and laboratory studies (Brown et al., 2000 and references therein; Jakobsen and Postma, 1999; Ludvigsen et al., 1998; Weiner et al., 1998). After day 42, SO_4 was exhausted and Fe(III) reducing conditions were likely to be dominating the system. In fact, concomitant occurrence of methanogenesis could not be excluded after day 42, favored by the slow rate of Fe(III)-reduction. Coupling the two processes with siderite precipitation, this could represent another limiting factor for the increase of Fe(II) concentration. Similar scenarios have been already reported in literature (Jakobsen and Cold, 2007). Additional potential explanations for the low Fe(II) concentration detected, aside from the fact that the investment of some DOC in biomass

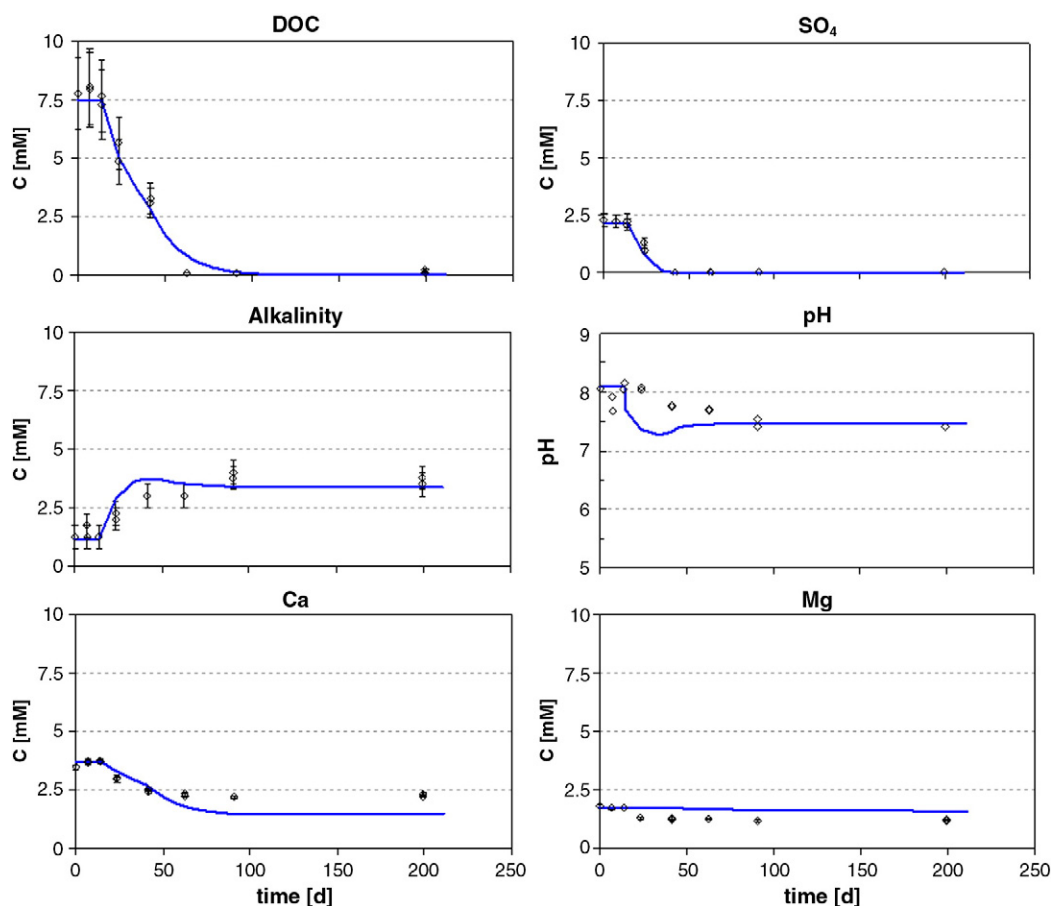


Fig. 6. Chemical evolution with time in the Fe(III)-reducing experiment: simulation "A" (solid lines) versus experimental data (dots).

growth implies a smaller Fe(II) production than mineralization, could be some Fe(II) adsorption onto clay surfaces (Wang and Van Cappellen, 1996) and/or its incorporation with some remaining solid Fe(III) to form magnetite (mixed Fe(II)–Fe(III) oxide) (Broholm and Arvin, 2000 and references therein; Brown et al., 2000; Lovley and Phillips, 1988).

Small crystals of calcite, Mg–Ca carbonates and siderite were observed on the surface of sediment grains (Fig. 2D), confirming that indeed precipitation of carbonates was limiting the increase of alkalinity.

When modeling the experiment with CHEPROO, the best fits of experimental data were obtained under the hypothesis of organic matter being degraded during the first part of the experiment by SO_4 and, to a lesser extent, by Fe(III); then, after SO_4 depletion, concomitant Fe-reduction and methanogenesis were assumed to be responsible of the organic substrate consumption (Fig. 6). The most important parameters used in the simulation, identified as simulation “A” in the following, are detailed in Table S5 of the Supporting Information.

Among the numberless combinations of hypotheses likely accounting for the complex geochemical evolution of the experiment, the simulations presented in this section were carried out without taking into account biomass growth during SO_4 -reduction and methanogenesis.

The need of considering the occurrence of Fe(III)-reduction and methanogenesis beside SO_4 -reduction was suggested by the bigger departure between model results and measurements (Fig. 7) when considering the degradation of organic matter coupled with one of the following processes: SO_4 -reduction (in this case biomass growth was included) (simulation “B”); SO_4 -reduction and slow Fe(III)-reduction (simulation “C”); SO_4 -reduction and fast Fe(III)-reduction (simulation “D”); and SO_4 -reduction and methanogenesis (inhibited by SO_4 presence) (simulation “E”).

To be noted, moreover, that simulation “A” was obtained by considering inhibition of Fe(III)-reduction in the presence of SO_4 . On the contrary, again, computed DOC and Alkalinity showed worst fits to experimental data (results not shown).

3.1.4. SO_4 -reducing experiment

Results from the sulfate reducing experiment are shown in Fig. 8. A small decrease of DOC occurred prior to day 7, part of it (a maximum of 0.2 mmol L^{-1}) being associated with the depletion of the initial 0.1 mmol L^{-1} of NO_3 (not shown). Nevertheless, the significant decreases in both DOC and SO_4 could be observed between days 18 and 65. By then, the initial 10.2 mmol L^{-1} of DOC have been almost depleted (about 0.3 mmol L^{-1} remaining) and SO_4 concentration has decreased from the initial 5.3 mmol L^{-1} to 2.1 mmol L^{-1} . After day 65, SO_4 continued decreasing down to 1.3 mmol L^{-1} at the end of the experiment despite the fact that DOC had been already practically exhausted. In parallel, alkalinity increased continuously from the initial 1.5 mmol L^{-1} to 4.3 mmol L^{-1} at day 65, and up to 6.3 mmol L^{-1} by the end of the experiment. Ca and Mg concentrations decreased from 3.6 to 1.3 mmol L^{-1} and from 1.8 to 0.9 mmol L^{-1} , respectively. A very low concentration of Mn, never exceeding 0.01 mmol L^{-1} , was detected during the whole experiment, probably associated to dissolution or reduction of some Mn mineral naturally present in the sediments, this having a negligible effect on DOC.

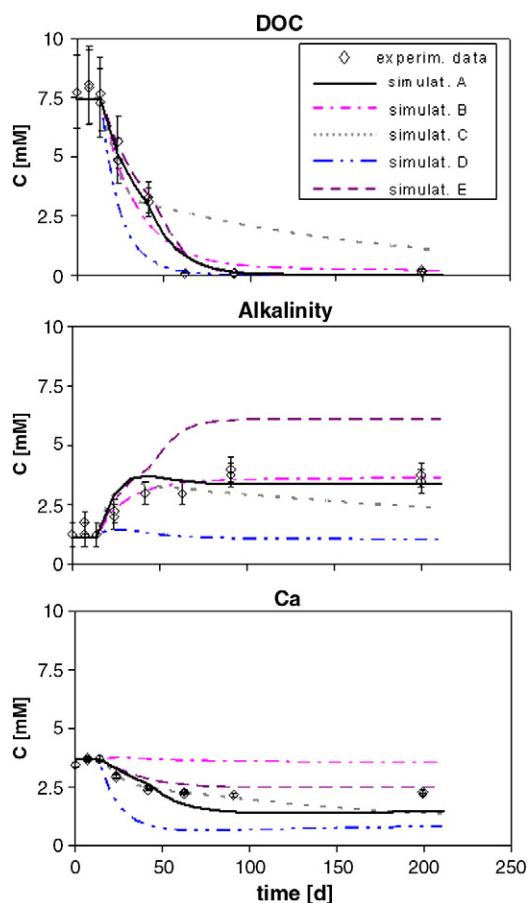


Fig. 7. Chemical evolution with time in the Fe(III)-reducing experiment: simulations (lines) versus experimental data (dots). Namely, regarding simulations: (A) SO_4 -reduction, Fe(III)-reduction and methanogenesis (both the latter inhibited by SO_4), (B) SO_4 -reduction (biomass growth included), (C) SO_4 -reduction and slow Fe(III)-reduction, (D) SO_4 -reduction and fast Fe(III)-reduction, (E) SO_4 -reduction and methanogenesis (inhibited by SO_4).

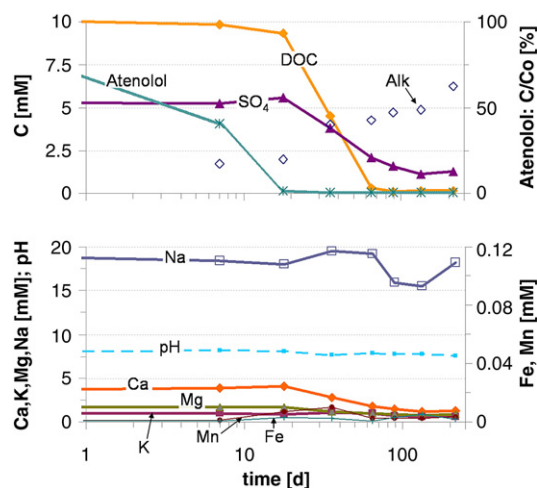
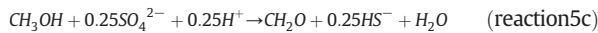
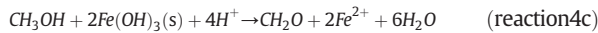


Fig. 8. Chemical evolution with time in the SO_4 -reducing experiment.

As explained in the case of Fe(III) and as result of SO_4 -reduction, two alternative hypotheses might be formulated: (1) HS^- remained in solution, or (2) HS^- precipitated as FeS with the Fe^{2+} produced by reduction of some of the Fe(III) oxides present in the original sediment. Hypothesis (1) was supported by a decrease in charge balance error for the water samples up to day 65. Hypothesis (2) was supported by the dark color in the sediments after day 18.

Assuming hypothesis (1), by using the actual decrease in sulphate concentration (3.2 mmol L^{-1}) and DOC (9.9 mmol L^{-1}), the stoichiometries in Table 1 and the additional reactions:



to take into account that in the case of methanol the formation of biomass (simplified formula: CH_2O) requires a partial oxidation too, it could be calculated that a total amount of 5.3 mmol L^{-1} DOC was mineralized up to day 65 while 4.6 mmol L^{-1} were inverted into microorganisms' growth (47%). Since sulfate still decreases and alkalinity increases after DOC is nearly exhausted, we concluded that biomass was reused (Alexander, 1999). Under this assumption and according with stoichiometry,

to reduce the remaining $0.8 \text{ mmol L}^{-1} \text{SO}_4$, 0.17 mmol L^{-1} of the remaining DOC and some 1.5 mmol L^{-1} of accumulated biomass were further mineralized after day 65. Thus, the "net" investment of DOC into biomass during the whole experiment amounted to 3.1 mmol L^{-1} . Global balance of the experiment implies that 69% of organic carbon decay was associated to mineralization of the organic substrates coupled with sulfate reduction, and 31% was inverted into microorganisms' growth.

Assuming hypothesis (2), with some concomitant Fe(III) reduction occurring, a similar calculation could be made and the global balance resulted in 69 to 72% of organic carbon diminution associated to substrates mineralization coupled with sulfate reduction, a 7 to 10% coupled with Fe(III) reduction, and 21% of organic carbon inverted into microorganisms' growth.

Most likely, actual processes lied between these two extreme cases, implying that a fraction of the HS^- remained in solution as aqueous species while part of it was precipitated as FeS. In either case, SO_4 -reduction coupled with organic matter degradation was the dominating process during the experiment. A microbial equilibration period or the enmasking of SO_4 -reduction early stages by the analytical errors could explain the insignificant changes characterizing water chemistry during the first 7 to 18 days of the experiment.

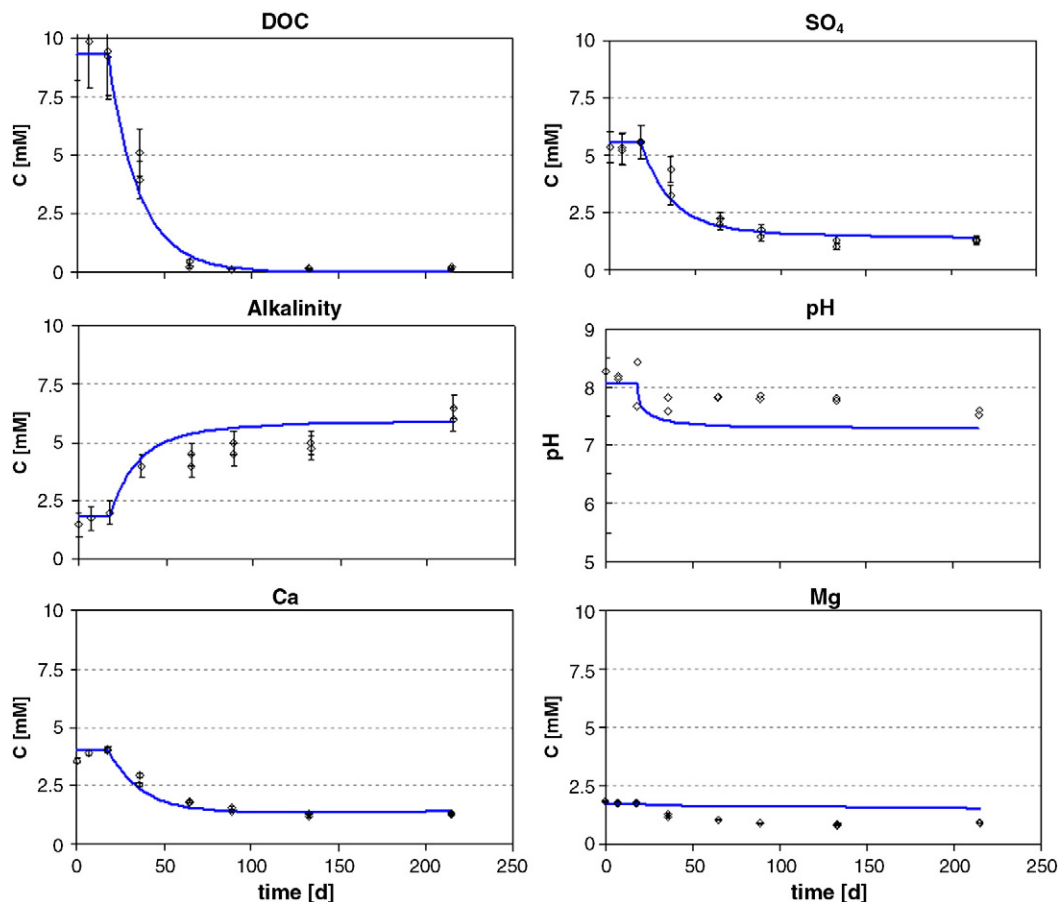


Fig. 9. Chemical evolution with time in the SO_4 -reducing experiment: simulations (solid lines) versus experimental data (dots).

The total DOC mineralization estimated under both hypotheses (1) and (2) (6.9 and 7.9 mmol L^{-1} , respectively) were higher than the measured increase in alkalinity (4.8 mmol L^{-1}). Attributing the net reduction of Ca and Mg concentrations (2.3 mmol L^{-1} and 1 mmol L^{-1} , respectively) to the precipitation of Mg and Ca carbonates, and taking into account the inorganic carbon present as gaseous phase in the headspace of the bottles as well as the precision of alkalinity measurements, the expected alkalinity fitted quite well the measured value. The resulting error in the global carbon balance amounts to about 13% and 1% under hypothesis (1) and (2), respectively. Indeed, Mg and Ca carbonate crystals were observed on the sediment grain surfaces.

The results for the simulations of the experiment chemical evolution under an intermediate case between hypothesis (1) and (2) are reported in Fig. 9. The most important parameters used are detailed in Table S6 of the Supporting Information.

3.1.5. “natural conditions” experiment

Results from the “natural conditions” experiment are shown in Fig. 10. The initial 2.5 mmol L^{-1} DOC were almost completely depleted, starting from the very beginning of the experiment. The initial 0.2 mmol L^{-1} of dissolved O_2 (not shown) and 0.1 mmol L^{-1} of NO_3 were totally removed after 1 and 3 days, respectively. Dissolved Mn was observed at approximately day 10, reaching a maximum concentration of 0.03 mmol L^{-1} at day 62; then, Mn concentration decreased to about zero at the last sampling time of 192 days. Some dissolved Fe was detected between days 15 and 62, never exceeding 0.02 mmol L^{-1} . The initial 2.3 mmol L^{-1} SO_4 decreased with some fluctuations to 1.9 mmol L^{-1} , from day 15 to the end of the experiment. The alkalinity showed an overall increase during the experiment, from 0.5 to 2.3 mmol L^{-1} .

To sum up, since no specific redox state was deliberately stimulated in the “natural conditions” experiment, the organic matter degradation reactions occurred in the expected se-

quence (Table 1, set “a” of reactions), until complete depletion of the specific electron acceptor (e.g. oxygen and nitrate) or, finally, of the electron donor. Aerobic degradation dominated the first day, and nitrate reduction appeared to control degradation until day 3. From there on, Fe and Mn-reducing conditions were found. These overlap with the SO_4 -reduction, which occurred after day 15. The presence of small zones of dark color in some of the retrieved solid suggested that some precipitation of iron sulfide occurred. Also the decrease in Mg (0.4 mmol L^{-1}), Mn (0.03 mmol L^{-1}) and Fe (0.02 mmol L^{-1}) suggested that some carbonate precipitates. The overall Ca increase of about 0.8 mmol L^{-1} suggested carbonate dissolution, even if its concentration has been fluctuating during the experiment. Small crystals of Mg–Ca carbonates and siderite have been observed on the surface of sediment grains.

3.2. Fate of atenolol

As example of application of the described microcosm study to the fate of emerging organic micropollutants under different redox conditions, the results for the β -blocker atenolol are presented. The temporal evolution of its average normalized concentration (with respect to each actual initial concentration C_0) for the 6 sets of experiments described above is shown in Fig. 11. The error bars were calculated by taking into account the analytical errors and the difference between duplicate batches results. Concentrations are presented in relative terms, normalized as C/C_0 , in order to remove systematic errors from the analysis.

The behavior of atenolol was similar in the NO_3 -reducing experiment than in the first 10 days of the “natural conditions” experiment. Little removal of atenolol was observed until day 1.5. Then concentrations started to decrease, following almost the same trend in both experiments and reaching an overall removal of about 50% at day 10. In the Mn(III/IV)-, Fe(III)- and SO_4 -reducing tests, the lack of intermediate sampling points hindered the identification of atenolol behavior during the first week. Nevertheless, also in these three sets of experiments the same overall removal of atenolol of about 50% could be observed at day 7.

After days 7–10, different evolutions of the concentration curves of atenolol could be identified for each set of batches.

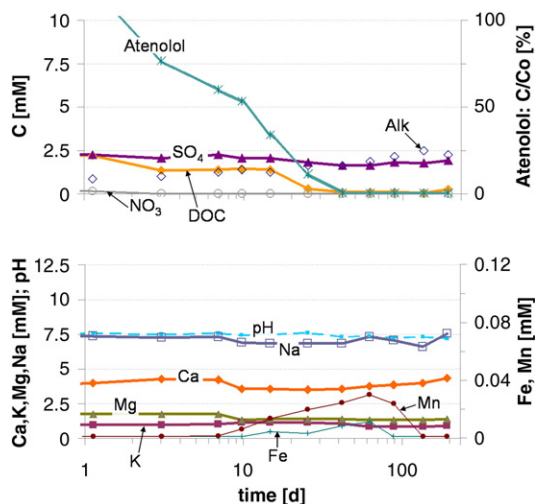


Fig. 10. Chemical evolution with time in the experiment performed under natural conditions.

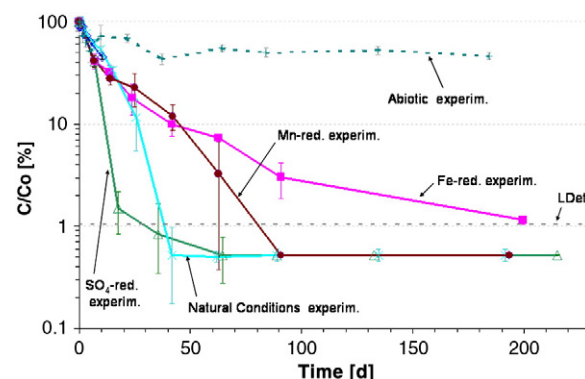


Fig. 11. Temporal evolution of the atenolol average normalized concentration in the different experiments. “LDet” stays for Limit of Determination.

At day 18, complete removal of atenolol was reached in the SO_4 -reducing experiment. In fact, through the evolution of the general hydrochemistry we could not confirm if during this period the target redox condition was still being established or had already developed. Under Mn-reducing conditions, 90% of removal was reached at day 42, up to complete removal at day 91. Similarly, in the Fe(III)-reducing experiment, under actual mixed Fe-/ SO_4 -reducing conditions, about 90% was removed at day 42. Complete depletion occurred later on, almost at the end of the experiment, under the sole Fe-reducing or mixed Fe-reducing/methanogenetic conditions. In the “natural conditions” experiment, under some mixed Mn-/Fe-/ SO_4 -reducing conditions, complete removal of atenolol was reached already at day 42.

Results for the abiotic experiment evidenced that, within the first 5 days, the removal trend for atenolol was the same observed in the NO_3 -reducing and “natural conditions” experiment, determining an overall removal of about 50%. After day 5, taking into account the error bars, no additional removal could be observed.

Thus, the initial ~50% removal of atenolol occurring within the first 5–10 days in all experiments could be attributed to some abiotic process, most likely sorption on the sediment grains. By then, some microbial processes were responsible of the remaining 50% removal of atenolol. Different evolutions could be observed depending on the experiment, i.e. mainly depending on the redox conditions dominating or being established in each system. The NO_3 -reducing experiment was too short to be compared with the remaining tests. Yet, among the latter, qualitatively it could be assessed that the faster atenolol biotic removal was observed in the SO_4 -reducing experiment, under the most reducing (already established or still being established) condition up to that moment.

Aiming at a better characterization of atenolol biotic removal, a number of samples from the different experiments were analyzed looking for the potential presence of transformation products, specifically for atenololic acid. This compound was identified as product of atenolol microbial hydrolysis in a study of Radjenovic et al. (2008). The expectation to find atenololic acid at least in the NO_3 -reducing experiment was fostered by its occurrence in a similar experiment we carried out with much higher (1 mg L^{-1}) initial atenolol concentration (Barbieri, 2011). Unfortunately, we found out that the analytical methodology used was not adequate to detect the potential presence of atenololic acid at concentrations in the ng L^{-1} order of magnitude (in the experiment the maximum attainable concentration was $1 \mu\text{g L}^{-1}$). Thus, its presence could not be confirmed due to analytical restraints.

4. Conclusions

The following concluding considerations and remarks can be made on the present study:

- The desired redox states have been quite successfully created and sustained in each experiment. It is worth pointing that the use of natural sources of Mn and Fe, as in the Mn- and Fe-reducing experiments, is realistic but complicates the development of controlled redox conditions. Natural sources are often quite crystalline, which

slows down dissolution to the point of making it the rate limiting process.

- The assessment of the dominating redox states has been achieved by a thorough monitoring of water chemistry, focused on the redox-sensitive species but including major and minor ions too. Precipitation/dissolution of minerals as well as biomass production have to be taken into account for a correct interpretation of the main processes involved. Inspection of the sediments from the disassembled batch experiments through SEM-EDS has been fruitful to confirm the occurrence of such processes.
- However, further improvements are required. Specifically, dissolved sulfide and methane should be analyzed to better assess sulfate reducing conditions, especially in its early stages, and to check possible occurrence of methanogenesis. Additional desorption experiments could confirm Mn^{2+} and Fe^{2+} adsorption onto clay surfaces and/or exopolymeric substances (EPS). As general rule, whenever possible, the evaluation of the microbial state during the experiments (e.g.: identification of microbial communities, measurements of hydrogen, etc.) would be also advisable as complementary tool for the identification of the prevailing redox state.
- Numerical modeling proved useful in confirming the concepts described above with literature kinetic rates. Matches between computations and observations could have been improved by varying the rates of carbonates precipitation, and by postulating likely occurring sorption onto biofilms. Departures between model results and measurements are small, but generally suggest an intricate coupling between biologic and inorganic processes.
- The sampling schedules have proven adequate for monitoring the temporal evolution of aqueous chemistry and micropollutant concentrations. Still, in the case of Mn-/Fe-/ SO_4 -reducing experiments some additional sampling point during the first week could have been useful to confirm atenolol early removal trends.
- One of the aims of the study was to test systems representative of real aquifers and of conditions occurring either naturally or possibly being stimulated during managed artificial recharge operations. Such conditions may vary spatially and temporally along with recharge cycles and recharge water composition. Thus, the microbial communities naturally existing in the sediments used in the experiments, which were expected to carry out the biodegradation of organic matter and the removal of micropollutants, were not previously adapted to the redox conditions of interest. As a consequence, the first part of each biotic experiment was characterized by a transition stage (of different duration) until the target redox state could be effectively established or observed. This hindered the interpretation of atenolol results. Anyway, after a common removal for all experiments, which we associate to abiotic processes, different microbial removal trends were observed for atenolol, depending on the different sets of batches, each one of them characterized by different redox conditions. The case of atenolol confirms that the redox state of the system could exert an influence on micropollutants behavior. Even if neither the NO_3 -reducing experiment was long enough to compare nitrate reducing conditions with the more reducing

systems nor exact patterns could be isolated for each specific redox state, the faster atenolol biotic removal rate was observed in the SO_4 -reducing experiment, under the most reducing condition (while being established or in its early stage) up to that moment.

- Actually, correctly identifying of the actual biotic processes responsible for the removal of micropollutants (i.e. to distinguish biotransformation from biodegradation or even mineralization) requires the use of specific techniques, such as the use of isotopically labeled compounds and/or the identification of already known/new transformation products. In our study, for the case of atenolol we sought for its transformation product atenololic acid, but the presence of such compound could not be confirmed due to analytical restraints.
- Due to design constraints, the concentration of the easily degradable organic substrates used in the experiments were higher than those naturally present in aquifer systems or in most recharge waters, which likely affected the growth of the microbial communities present in the microcosms. Thus, regarding the micropollutant atenolol the extrapolation of its biotic removal rates to natural subsurface environments would have to be faced carefully, being not straightforward. Still, the microcosm study proved the feasibility of specific redox environments to develop at test site and the capability of the local microorganisms to eliminate the target micropollutant, providing as well some overall removal pattern under the tested settings. In the end, such scenarios could eventually be promoted during artificial recharge at test site if less favorable removals of atenolol are observed under the spontaneously occurring conditions.
- The removal of atenolol reported in the literature varies between 0 and 60% in conventional wastewater treatments, improving up to 77% removal in advanced treatments such as Membrane Bioreactors (Gros et al., 2010; Radjenovic et al., 2009 and references therein). Thus, the overall complete removal observed in the experiments performed within this study suggests that the whole processes occurring in aquifers constitute a potentially efficient alternative water treatment for atenolol. Depending on the redox state naturally occurring or possibly being deliberately stimulated in field applications, the time needed for a complete removal may be ensured by the large residence times in aquifers. Actually, the results from the “natural conditions” experiment, which better resemble the potential conditions spontaneously occurring within the aquifer at Sant Vicenç test site during recharge, look very promising.

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