

Simulation of Water Circulation over a Model of a Submarine Canyon by Using FIC-FEM Numerical Model

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Abstract: A set of numerical model experiments has been conducted to simulate the circulation driven by oscillatory forcing over a theoretical continental slope configuration used previously in laboratory experiments. The test case considered was the numerical simulation of the flow over a model of a submarine canyon, and the numerical model used in the analysis was a coastal ocean model version based on an adaptation of the finite-calculus–finite-element method (FIC-FEM) approach implemented in the commercial package Tdyn. Two cases were analyzed involving changes in fluid density. Structured and unstructured finite-element spatial discretizations were generated for the same study domain to compare the resulting velocity field with outputs from the laboratory experiments and to assess which mesh provided a better representation of the complex geometry of the channel model and the water circulation process. The comparison between the laboratory results from the reference article and the output of the numerical model showed good agreement in the structure and magnitude of the phase-averaged and residual velocity fields.

Author keywords: Submarine canyon; Coastal flow; Tdyn; Finite element; FIC.

Introduction

The continental shelf is a transition zone between the coastline and the open sea, and is interrupted at irregular intervals by submarine canyons that are present on the shelf edge along most of the world's coastlines. Coastal canyons have long been of interest to scientists because of their geological and biological importance. Physical oceanographers find evidence that these topographic features are preferred pathways for shelf-slope transport of material and energy and play important roles in distributing suspended particulate matters and exporting sediments to the open ocean (Durrieu de Madron 1994). From a biological point of view, enhanced species diversity and biological productivity can be found in submarine canyons. Sardà et al. (1994) find evidence that these regions are preferential recruitment habitats, whereas Gili et al. (2000) see canyons as biodiversity hot spots.

Studies on the effects of canyon topography on the regional circulation are relatively few; they have been reviewed by Hickey (1995) and Huthnance (1995). Field observations are very rare

[e.g., Hickey (1997) and Bosley et al. (2004)] because physical measurements in submarine canyons are among the most complicated to make. Various factors contribute to the difficulty of performing field measurements including the steepness of canyon slopes and the maintenance of the moored arrays of conductivity-temperature-depth (CTD) sensors in water because of intense fishing activities concentrated in these areas. These few field studies helped explain some previously unknown processes, but also presented shortcomings that include high cost and the difficulty of attributing a specific cause to an observed effect because of the simultaneous occurrence of many physical processes. More recently, the focus of attention has been the use of numerical experiments to address these drawbacks. Several numerical experiments relating to submarine canyons have been conducted and have shown that circulation within the submarine canyons is characterized by both downwelling and upwelling events. Some numerical models (Hurlburt 1974; Peffley and O'Brien 1976) have demonstrated that canyons affect the patterns of regional upwelling and that upwelling is enhanced on the downstream side of a canyon. Klinck (1988) attempted to model the circulation in a submarine canyon and its interaction with shelf flow using a vertically walled channel, providing the first insight into canyon-flow interaction. His studies showed that the width of the canyon is an important parameter, and the interaction was found to be smaller for canyon widths smaller than the current width or the internal Rossby radius of deformation (ratio between the phase speeds of the long internal waves to the Coriolis parameter). Two models (Allen 1996; Klinck 1996) with realistically steep and abrupt topography provided major contributions to the understanding of the interaction of shelf flow with coastal canyons. The outputs from these models were qualitatively consistent with the set of spatially comprehensive observations available. Several attempts to use numerical modeling to simulate current circulation followed, and Haidvogel and Beckman (1995) conducted an intermodel comparison.

Numerical models, supported by adequate field data, provide increasingly realistic results that can be used to predict general flow patterns, but the absence of suitable validation tools leaves the issue

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of using numerical models with a few doubts. As an alternative, recently formulated numerical models have relied on comparisons with the outputs of other numerical and laboratory models for validation purposes even though comparisons between numerical models have been proven inconclusive.

In this paper, the effects of sea floor topography, particularly the presence of submarine canyons, on water circulation were analyzed through a theoretical continental slope configuration used previously in laboratory experiments. The plan was to compare the outputs of the numerical model used in the analysis with the results obtained from the laboratory experiment (Pérenne et al. 2001).

Description of the Numerical Model

The numerical model employed is a coastal ocean model version based on an adaptation of the finite-calculus–finite-element method (FIC-FEM) (Oñate and García-Espinosa 2001; García-Espinosa and Oñate 2003; Oñate et al. 2006) implemented in the commercial package Tdyn. The stabilized governing equations for the incompressible fluid and the free surface were derived by the use of a finite-calculus (FIC) procedure, which is a technique based on writing the different balance equations over a domain of finite size and retaining the higher order terms. The resulting stabilized numerical solution scheme, coupled with the use of finite-element discretization of the domain, allowed the numerical model greater flexibility, especially when dealing with the complex geometry of irregular coastlines and seabeds.

General FIC Equations for Viscous Incompressible Flow

The velocity and pressure fields of an incompressible fluid moving in a domain $\Omega \subset R_d (d = 2, 3)$ are described by the incompressible Navier-Stokes equation

$$\frac{\partial \rho u_i}{\partial t} + \frac{\partial}{\partial x_j} (\rho u_i u_j) + \frac{\partial p}{\partial x_i} - \frac{\partial s_{ij}}{\partial x_j} = \rho f_i \quad \frac{\partial u_i}{\partial x_i} = 0 \quad i, j = 1, d \quad (1)$$

where $1 \leq i, j \leq d$; ρ = density of the fluid; u_i = i component of the velocity field \mathbf{u} in the global reference x_i ; p = pressure field; and s_{ij} = viscous stress tensor defined by

$$s_{ij} = 2\nu \left(\varepsilon_{ij} - \frac{1}{3} \frac{\partial u_k}{\partial x_k} \delta_{ij} \right) \quad \varepsilon_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right) \quad (2)$$

The stabilized FIC form of the governing differential equations in Eq. (1) can be written as

$$r_{m_i} - \frac{1}{2} h_{ij}^m \frac{\partial r_{m_i}}{\partial x_j} = 0 \quad \text{in } \Omega \quad (3)$$

for momentum and

$$r_d - \frac{1}{2} h_j^d \frac{\partial r_d}{\partial x_j} = 0 \quad \text{in } \Omega \quad (4)$$

for mass balance. The terms r_{m_i} and r_d are the residuals of Eq. (1) and h_{ij}^m and h_j^d are the characteristic lengths that represent the dimensions of the finite domain in which the mass balance and the momentum are imposed. The details of the FIC method for the computation of the stabilization terms can be found in Oñate (1998).

The boundary conditions for the stabilized problem correspond to the following equations:

$$n_j \sigma_{ij} - t_i + \frac{1}{2} h_j n_j r_{m_i} = 0 \quad \text{on } \Gamma_t \quad u_j = u_j^p \quad \text{on } \Gamma_u \quad (5)$$

where n_j = vector component normal to the boundary; t_i and u_j^p = prescribed surface tractions and velocity; Γ_t and Γ_u = two sets of disjoint components of the boundary $\partial\Omega$ where the Neumann and Dirichlet boundary conditions for the velocity are prescribed; and σ_{ij} = total stress defined as

$$\sigma_{ij} = s_{ij} - p \delta_{ij} \quad (6)$$

Eqs. (3)–(5) are the starting points in the derivation of the stabilized finite-element method (FEM) to solve the incompressible Navier-Stokes equations. The equations are multiplied by generic weighting functions and the resulting weighted momentum and mass balance equations are integrated over the domain using the finite-element method (García-Espinosa et al. 2008).

An interesting feature of the FIC formulation is that it allows the use of equal order interpolation for the velocity and pressure variables (Oñate et al. 2004).

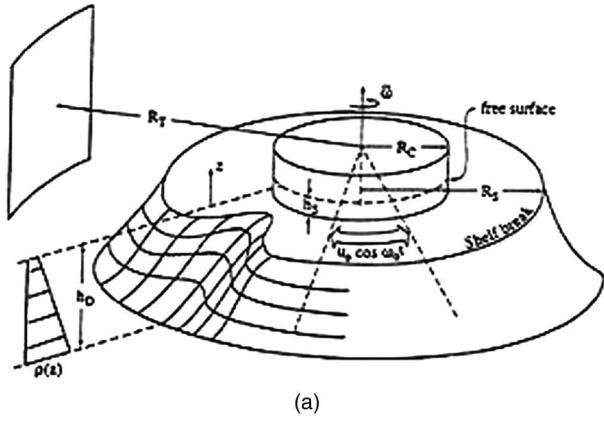
Physical Model Geometry

The numerical experiments were conducted in a three-dimensional representation of the domain of study configured to match the physical system setup for the laboratory experiment performed by Boyer et al. (2000). It has a circular tank of radius $R_T = 0.9$ m and a model of a circular continental shelf geometry [Fig. 1(a)]. The topographic model of the circular continental shelf was positioned in the center of the tank and consists of an annular coast (vertical), shelf (horizontal), and continental slope interrupted by a submarine canyon [Fig. 1(b)]. The horizontal shelf begins along a vertical coastline of radius R_C and the shelf break is located at radius R_S . The canyon incises both the continental shelf and the offshore slope. Its length is $2L$, measured along the axis from the point at which it intersects the horizontal shelf to the radius at which the sloping region reaches the deepwater floor. Its characteristic width W is the distance along the shelf break between the points at which the canyon intersects the shelf. The deep ocean is the part between the continental slope and the outside wall of the test cell.

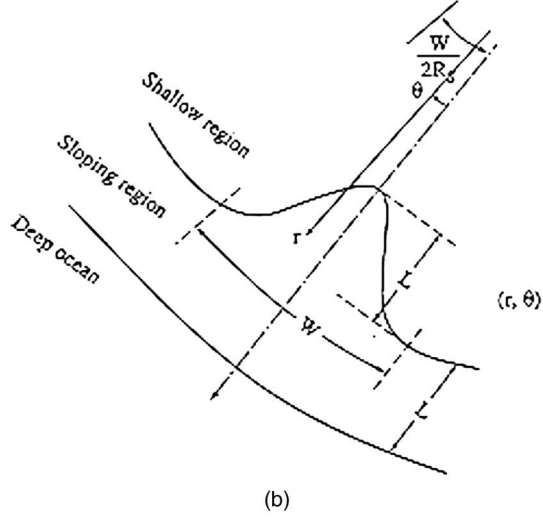
The tank was filled with fluid, and both homogenous and linearly stratified density situations were considered. The objective was to find the effect of fluid density—aside from the presence of a submarine canyon—on circulation. The system was made to rotate and the mean rotation of the turntable was not varied during the experiment. The background rotation of the tank was set at a constant Coriolis parameter of $f = 0.5 \text{ s}^{-1}$, and the simulated oceanlike flow was the along-isobath oscillatory flow defined as

$$v_\theta = -u_0 \frac{r}{R_S} \sin(\omega_0 t) \quad (7)$$

where $\mathbf{u}_0 = 1 \text{ cm s}^{-1}$ is the amplitude of the velocity at the shelf break, which is located at $r = R_S$; ω_0 = oscillation frequency; and t = time. The rotation was modulated sinusoidally to establish an oscillatory, along-shelf, background motion relative to an observer fixed to the canyon. Table 1 contains the values of the dimensional parameters.



(a)



(b)

Fig. 1. Physical and numerical models: (a) schematic diagram and (b) divisions of the continental shelf model (P  renne et al. 2001, with permission)

Table 1. Dimensional Values for the Numerical Model

Parameter	Value
R_C (radius of the coast)	35 cm
R_S (radius of the shelf-break)	55 cm
R_T (radius of the test tank)	90 cm
h_s (depth over the shelf)	2.5 cm
h_D (maximum depth)	12.5 cm
W (width of the canyon)	20 cm
L (length of the canyon)	15 cm
f (Coriolis parameter)	0.5 s^{-1}
ω_0 (forcing frequency)	0.26 s^{-1}
k_h (horizontal coefficient of diffusivity)	$1.5 \text{ m}^2 \text{ s}^{-1}$
K_v (vertical coefficient of diffusivity)	$0.1 \text{ m}^2 \text{ s}^{-1}$
U_0 (alongshore velocity forced at the shelf break)	1.0 cm s^{-1}
R_{ot} (temporary Rossby number)	0.52

Computational Mesh and Parameters

Given the complicated geometry of the study domain, the elemental partitioning was performed by using two different spatial discretization approaches. Initially, to compare the results with the studies of P  renne et al. (2001), the volume domain was discretized by using hexahedral elements in which the meshing was simplified by restricting it to a structured discretization (i.e., the number of

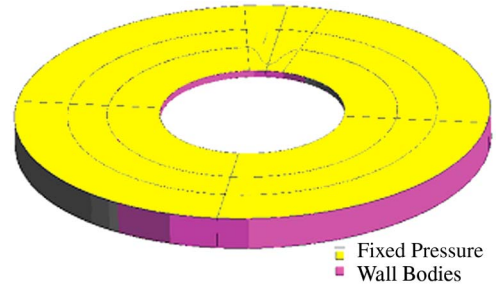
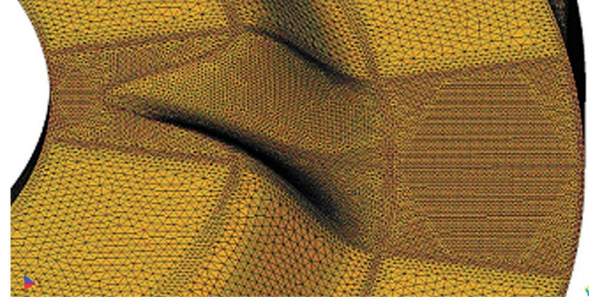
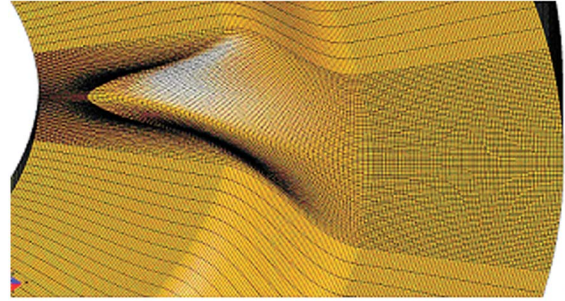


Fig. 2. Boundary conditions of the problem



(a)



(b)

Fig. 3. Canyon zone discretized using (a) unstructured and (b) structured mesh

vertical elements does not vary horizontally). The generated structured mesh has 166,980 nodes and 192,500 hexahedral elements.

The first computational mesh represented the domain of study well but was not optimized for the problem. Therefore, an unstructured mesh was generated that allows control of the size of the elements in the zone of interest. The resulting mesh has 2,323,547 tetrahedral elements with 395,219 nodes. In both cases, a Dirichlet boundary condition (Fig. 2) was imposed on the sides and bottom of the domain and a Neumann condition with null stress was imposed on the free surface.

Two types of spatial discretizations have been presented, and to discover which of the two types of mesh is more advantageous in terms of representing processes of interest, a new structured mesh was generated in which the characteristic length of the elements in the canyon zone approximates the size of the elements in the unstructured mesh. Fig. 3 shows a comparison of the canyon zone discretized using the two types of mesh generation.

Experimental Results

This case study analyzes the temporal development of a horizontal velocity field in the canyon zone. The data obtained in the present experiments were the phase-average flow and the residual current at

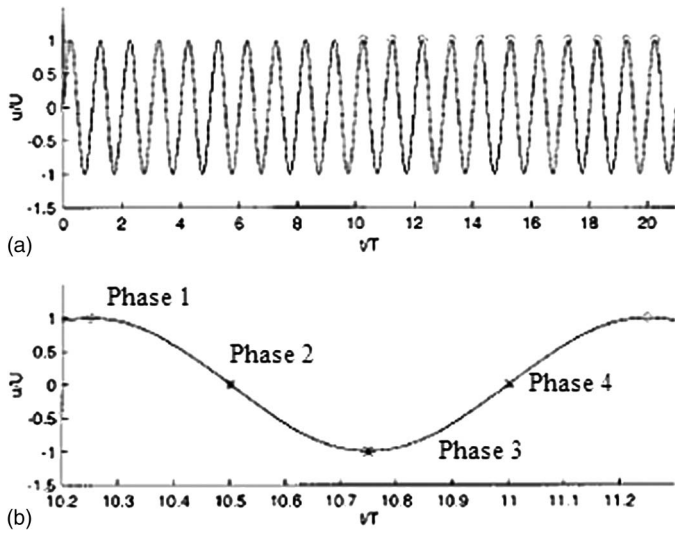


Fig. 4. Temporal averaging of the instantaneous velocity: (a) forcing sequence of the velocity that would be forced at the shelf break in the absence of a canyon and (b) definition of the four phases considered in the study (Pérenne et al. 2001, with permission)

the shelf break level ($z = -2.5$ cm). Fig. 4 defines the different kinds of temporal averages performed during the processing of the velocity data. The temporal averaging of the velocity field was calculated from the measurements involving cycles 11–20 after the start of the forcing; the phase averaging was thus performed over 11 cycles. The four phases considered are when the offshore flow is at the maximum in the downstream direction (Phase 1) and the succeeding quarters of a forcing cycle in which the downstream direction is in the right side of the canyon axis. The mean time of all the instantaneous velocity fields measured between the beginning

of the 11th cycle and the end of the 20th cycle is the experimental residual velocity field.

Two situations involving changes in the fluid density were considered to illustrate the effect of submarine canyons on coastal current circulation. The first case involves a homogeneous fluid density and the second considers a stratified fluid condition. Each of these cases was described by the horizontal velocity at three levels: mid-depth on the shelf, $z_1 = -h_S/2$; shelf break, $z_2 = -h_S$; and below the shelf break, $z_3 = h_S - (h_D - h_S)/4$.

The experiments were performed with $T = 24$ s as the oscillation period computed by using a temporal Rossby number defined as

$$Ro_t = \frac{2\pi}{Tf} \quad (8)$$

where f = Coriolis parameter; and Ro_t = temporal Rossby number. The time interval used for the simulation in all cases is $\Delta t = 1$ s with 504 time steps. A small region surrounding the canyon is used to present the simulation results.

Homogenous Fluid Case

The first test case, using a homogeneous fluid with a density of $1,000 \text{ kg/m}^3$, was the standard scenario for the flow pattern of the four phases defined in Fig. 4. Phase 1 exhibits the maximum rightward pattern, whereas phase 3 represents the maximum leftward flow phase. Phases 2 and 4 are for cases in which the background velocity is zero and the impending motion is in the opposite direction.

The simulation results obtained for the mid-depth on the shelf ($z_1 = -1.25$ cm) indicated that circulation at this level did not react strongly to the canyon bathymetry. It was weak over the center of the canyon, but a gradual increase in the magnitude of the velocity was observed at the shelf break. The velocity field displayed a

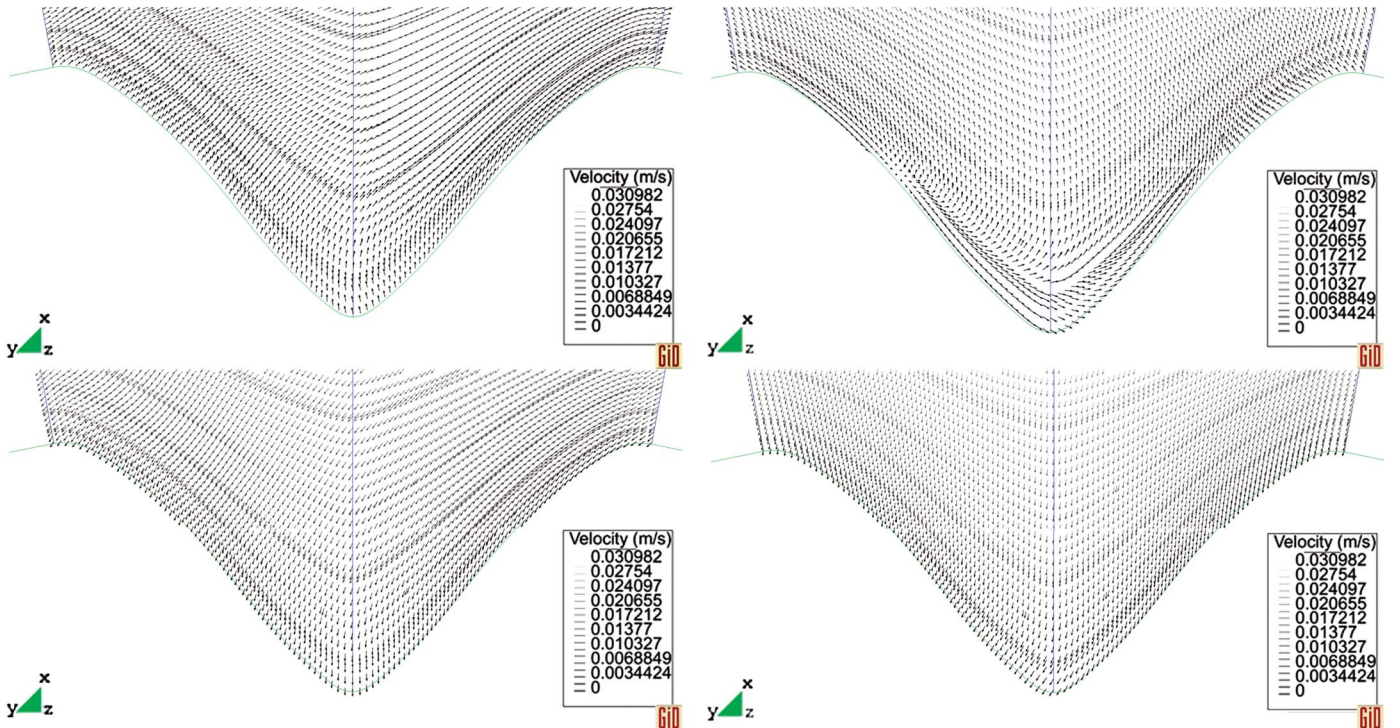


Fig. 5. Numerical model phase-averaged velocity fields obtained at $z = -5$ cm (structured mesh) for the four phases indicated in Fig. 4

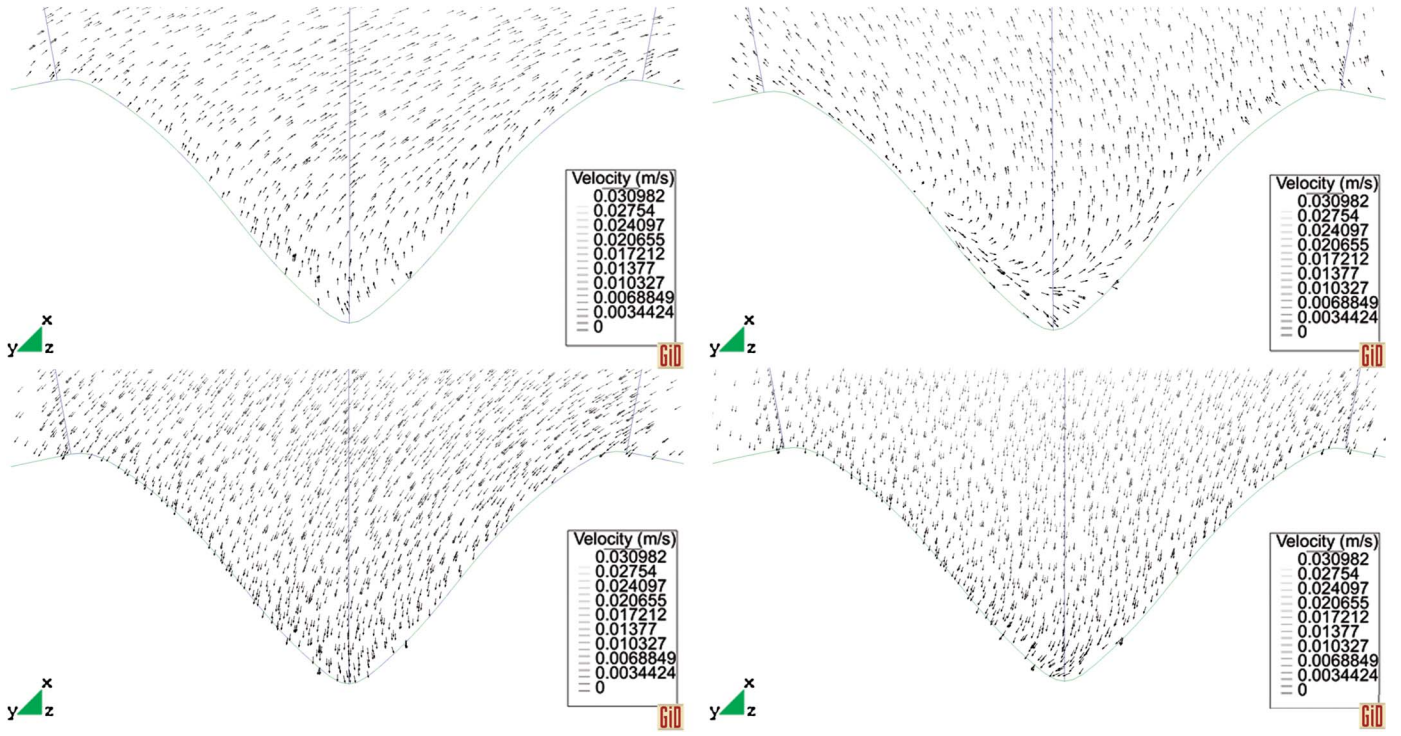


Fig. 6. Numerical model phase-averaged velocity fields obtained at $z = -5$ cm (unstructured mesh) for the four phases indicated in Fig. 4

different character when $z_2 = -2.5$ cm and the presence of the submarine canyon started to show a more noticeable effect on the imposed flow. The magnitude of the velocity field around the rim was about twice the horizontal circulation over the center of the canyon, which remained weak while those beyond the shelf break were not affected. Currents below the shelf break ($z_3 = -5$ cm) had similar characteristics as those above, with the current around the rim stronger than those within the center of the canyon (Figs. 5 and 6). Although both numerical results obtained from the two types of mesh generation had similar magnitudes, there were differences in the details; for instance, the results obtained from the unstructured mesh presented a clearer resolution of the ongoing process.

Stratified Fluid Case

The second case was a repeat of the previous, except that stable fluid stratification was now considered. The profile of density distribution corresponds to the value of the typical Brunt-Väisälä frequency $N = 2.5 \text{ s}^{-1}$ defined by the following equation (Dyke 2007):

$$N^2 = -\frac{g}{\rho} \frac{\partial \rho}{\partial z} \quad (9)$$

where g = local acceleration due to gravity; ρ = fluid density; and z = geometric depth.

To evaluate the performance of Tdyn, the simulation results were compared with the outputs of the laboratory experiments presented by Pérenne et al. (2001). The numerical grid used by Pérenne et al. has 250,000 nodes.

Fig. 7 shows the phase-averaged velocity fields obtained from the laboratory experiments, whereas Fig. 8 illustrates the numerical results from Tdyn. A good agreement in the direction of the numerical results and the laboratory measurement was noticed in phases 1 and 3 in which both results showed the maximum

rightward and leftward flow patterns. The development of a cyclonic eddy structure in phase 2 was not apparent in the numerical results compared to the laboratory version, although the velocity fields gradually changed direction from left to right at the head of the canyon to right to left near its mouth. The onshore current from the deepwater into the canyon along the canyon axis in phase 4 was more evident in the numerical model than its laboratory counterpart.

The residual currents obtained from the numerical experiments and the laboratory measurements are given in Fig. 9. The numerical

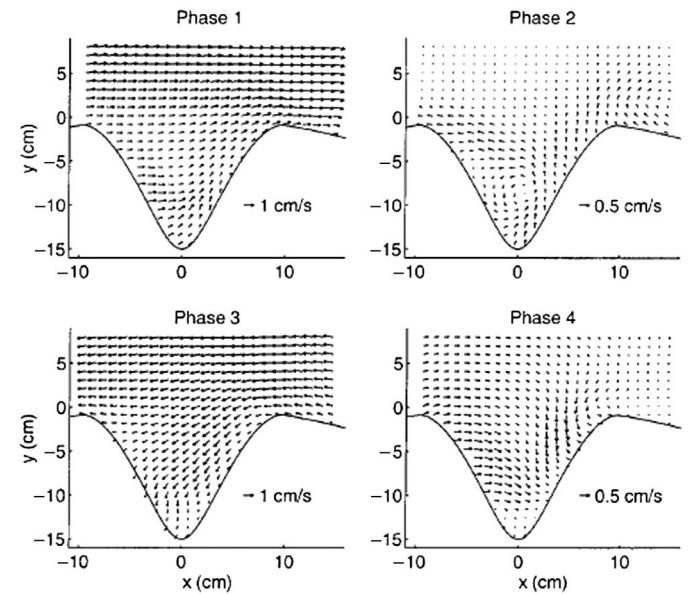


Fig. 7. Laboratory model phase-averaged velocity fields at $z = -2.5$ cm (Pérenne et al. 2001, with permission)

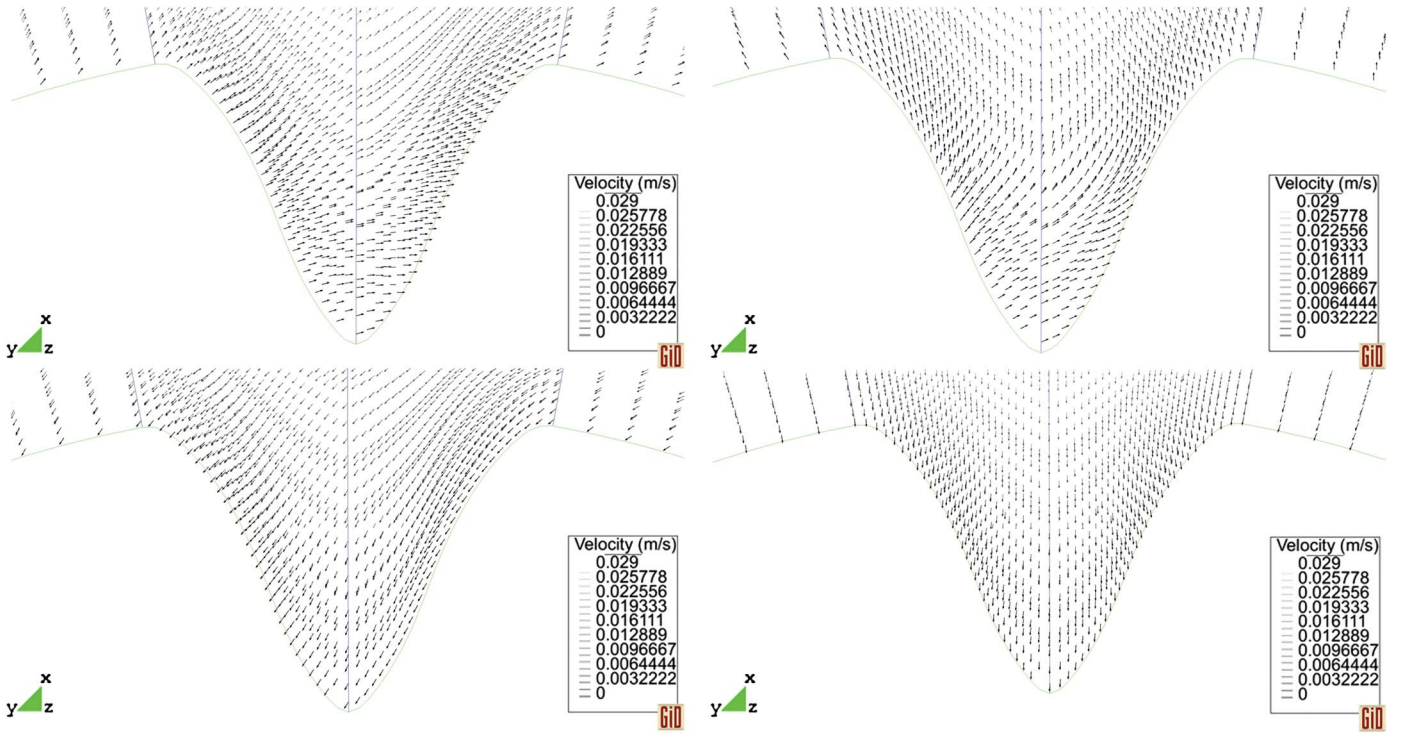


Fig. 8. Numerical model phase-averaged velocity fields at $z = -2.5$ cm

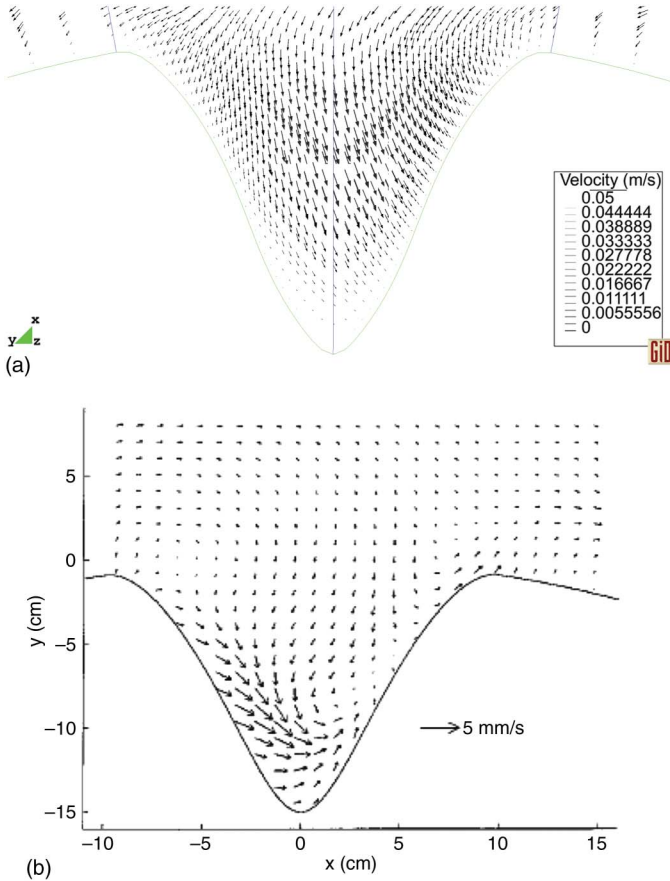


Fig. 9. Comparison of residual velocity fields at $z = -2.5$ cm: (a) numerical model and (b) laboratory model (P  renne et al. 2001, with permission)

results exhibited an intensification of the residual currents but were not confined to the vicinity of the upstream wall and spread up to the center of the canyon. The anticyclonic loop that can be found in the upstream wall of the canyon was clearly seen in the laboratory model but was absent in the numerical results. The difference in the qualitative agreement between the two results may be attributed to the different set of values for the vertical ($0.1 \text{ cm}^2 \text{ s}^{-1}$) and horizontal ($1.5 \text{ cm}^2 \text{ s}^{-1}$) coefficients of diffusivity employed in the numerical model, which was because of computational stability constraints.

Similar to the nonstratified simulations, model runs of the stratified situations using the structured and unstructured grids were performed. As mentioned previously, the second structured mesh [Fig. 3(b)] was generated with the characteristic length of the elements in the canyon zone similar in size to the elements in the same location in the unstructured mesh [Fig. 3(a)].

Fig. 10 shows phase 2 of the phase-averaged velocity fields at depths $z_2 = -2.5$ cm and $z_3 = -5$ cm. The general flow pattern was similar to the previous nonstratified case and the numerical results showed that the horizontal circulation in the center of the canyon was weak. A gradual intensification of the velocity fields was noted near the mouth of the canyon, but it did not move deeper inside the canyon. The circulation below the shelf, $z_3 = -5$ cm, showed similar characteristics as that above the shelf, but the presence of a cyclonic eddy structure was visible in the numerical results obtained from the unstructured mesh.

The difference between the stratified fluid condition and the previous case in which the fluid density was homogeneous can be better visualized in the magnitude of the velocity fields. Observation of the numerical results showed similarity in the general flow behavior; however, the increase in the fluid density reduced the magnitude of the velocity fields (cf. maximum value for homogeneous and stratified fluid condition) in the canyon zone (Fig. 11).

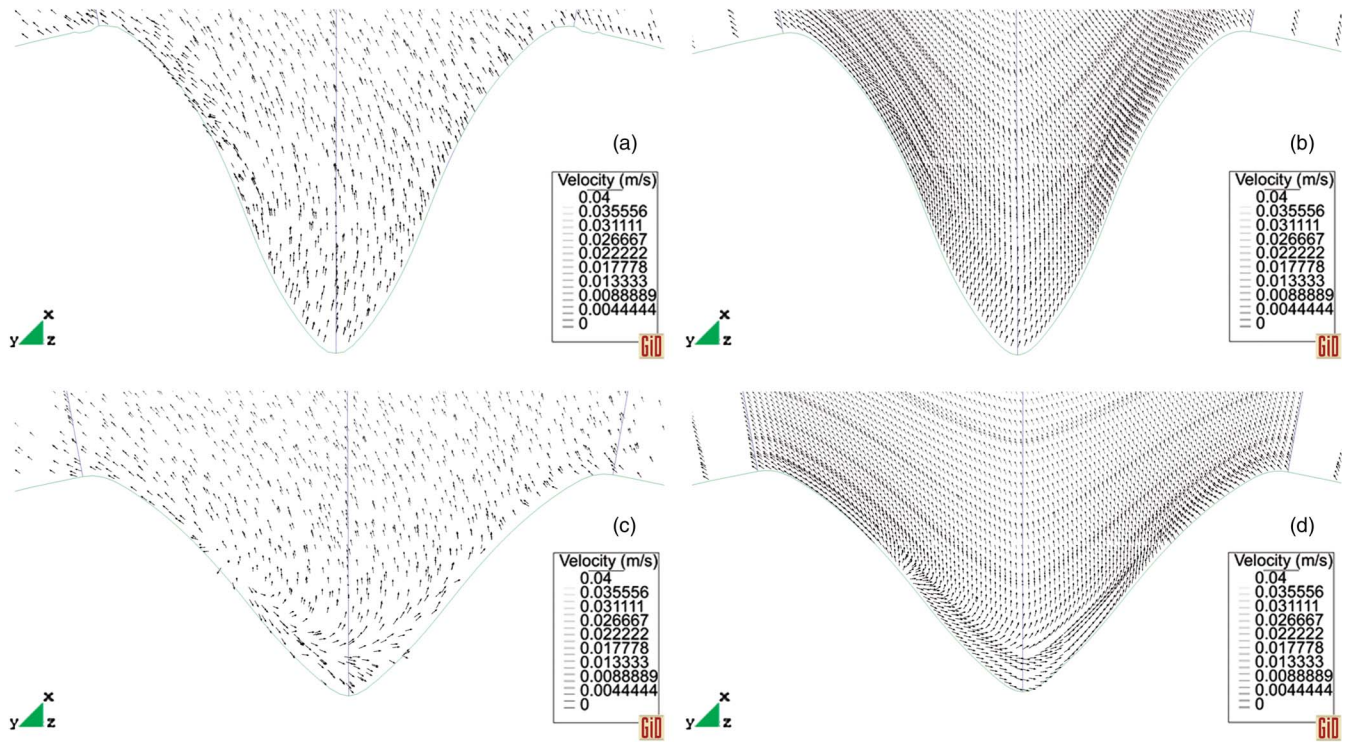


Fig. 10. Comparison of phase-averaged velocity fields (phase 2) (a) unstructured and (b) structured mesh at $z = -2.5$ cm; (c) unstructured and (d) structured mesh at $z = -5$ cm

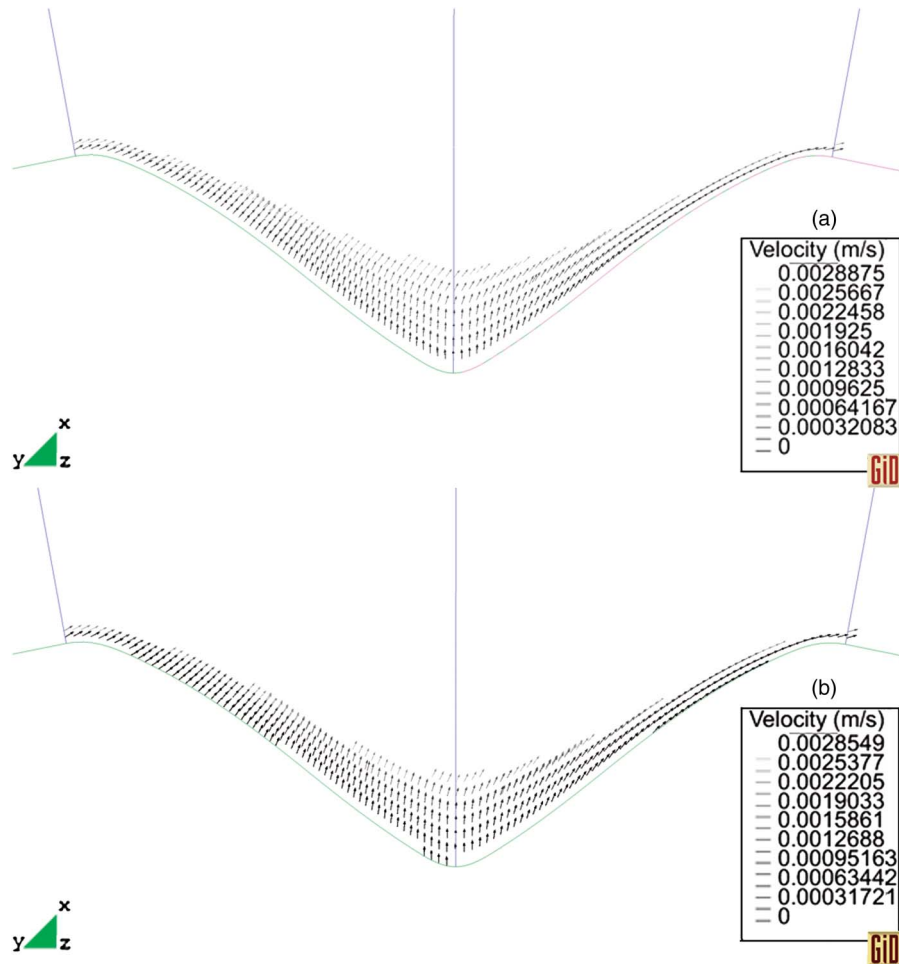


Fig. 11. Comparison of (a) homogeneous and (b) stratified fluid condition phase-averaged velocity fields (phase 1) at $z = -7.5$ cm

Conclusion

Several numerical experiments have been performed to simulate currents near and within a model of a submarine canyon located on a continuous shelf, shelf break, and continental slope topography. The model setup was taken from the laboratory experiments of Boyer et al. (2000) and the numerical model simulations were on a physical model scale. The dynamic setting involved an oscillatory background flow produced by modulating the rotation rate of the turntable. On the basis of model results, some issues regarding coastal flow near a submarine canyon were analyzed.

The results of this study showed that in the upper layers, the influence of canyon topography on current circulation was minimal. In deeper layers, the presence of submarine canyons changed the direction of the current circulation.

The effect of fluid density also contributed to limit the influence of the canyon on the overlying flow. The magnitude of the velocity fields in stratified conditions differed from their homogeneous counterparts because vertical motions were diminished in the former. This effect can be seen in the difference in the value of numerical results obtained from the homogeneous and stratified fluid conditions. The magnitude of the velocity fields obtained in the homogenous fluid condition was slightly stronger (1.15%) than the results obtained from stratified condition. The numerical results were in consonance with the observation that as the stratification increases, the effect of the canyon decreases (Klinck 1996).

The computation of phase-averaged and residual velocity fields generally agreed with the previous studies done by Pérenne et al. (2001). This suggests that the numerical model Tdyn can be used in a comprehensive canyon-flow analysis. Future lines of investigation include a real life application of the numerical model to study coastal flow circulation in Blanes Canyon (Flexas et al. 2008). Finally, the use of unstructured mesh provides better resolution in showing some processes that can not be seen in the use of structured grids.

Notation

The following symbols are used in this paper:

d = number of space dimensions;
 f = Coriolis parameter;
 h_D = maximum depth;
 h_j^d, h_{ij}^m = characteristic length distances;
 h_s = depth over the shelf;
 i, j = indices;
 L = length of the canyon;
 N = Brunt-Väisälä frequency;
 N_d = number of space dimensions;
 \mathbf{n} = unit outward normal to the boundary;
 n_j = vector component normal to the boundary;
 p = pressure field;
 R = domain of study;
 R_c = radius of the coast;
 R_{Or} = temporal Rossby number;
 R_s = radius at the shelf break;
 R_T = radius of the test tank;
 r = radius;
 r_{m_i}, r_d, r_ψ = residuals;
 s_{ij} = viscous stress tensor;
 T = forcing timescale;
 t = time parameter;
 t_i = prescribed surface traction;
 \mathbf{u} = velocity field;

u_i, u_j = i th and j th components of the velocity field \mathbf{u} ;
 u_o = amplitude of the velocity at the shelf break;
 u_j^p = prescribed displacement;
 ν = viscosity;
 W = width of the canyon;
 x = horizontal Cartesian coordinate along the x-axis;
 y = horizontal Cartesian coordinate along the y-axis;
 z = vertical Cartesian coordinate axis;
 Γ_t, Γ_u = components of the boundary;
 δ_{ij} = Kronecker's delta;
 ε_{ij} = strain rates;
 ρ = density;
 σ_{ij} = total stresses;
 Ω = analysis domain; and
 ω_0 = oscillation frequency.

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